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Proceedings A Natural and Anthropogenic Squeeze of the Monsoon Tidal Inlet ⁺

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Abstract: This study investigates the influences of natural and human-induced spatial blocking on 9 the morphological and hydrodynamic features of the Tien Chau tidal inlet. It is observed that the 10 absence of the downdrift barrier island and the appearance of the headlands form a squeeze tidal 11 inlet. Moreover, anthropogenic factors, including urbanization and aquacultures, block a large part 12 of the flood-tidal delta. These natural and human-induced spatial restrictions have significant influ-13 ences on the hydro and morphodynamic characteristics of the inlet. In order to obtain more insight, 14 a schematized numerical model was constructed in the Delft3D model and validated using meas-15 ured field data. 16

Keywords: squeeze; climate change; numerical model; field measurement; monsoon tidal inlet; human constructions

1. Introduction

Ky Lo river basin located in Phu Yen District, Quang Ngai Provinces, Vietnam, covers 21 approximately 1180 km². The Ky Lo river basin reaches the East Sea in the south region of 22 the Xuan Dai Bay through a narrow inlet named Tien Chau tidal inlet. This tidal inlet 23 environment was the study area (Figure 1). Xuan Dai Bay is bound by relatively steep, 24 sandy shorelines and has a large entrance with free exchange to the Southeast sea. Because 25 of the significant water exchange during the tidal cycle, The Xuan Dai Bay can be consid-26 ered tide-dominated, even it is located on a low meso tidal coast [4]. Xuan Dai Bay can be 27 divided into two major regions associated with the two prominent headlands. The north 28 headland mostly shelters the northside area of the Xuan Dai Bay, and the southeast head-29 land covers the southside area of the Xuan Dai Bay (Figure 2). The north headland pro-30 vides a natural protective barrier and blocks incoming waves from the north and east 31 directions. Thus, tidal processes tend to dominate the former region. A southeast head-32 land provides shelter from waves that are coming from the east and north direction. The 33 Tien Chau tidal inlet is located in the south area of Xuan Dai Bay. Most incoming waves 34 at this location are from the southeast direction. Therefore, a mix of tide and wave pro-35 cesses tend to dominate this area. 36

Situated at the South of Xuan Dai Bay, the Tien Chau tidal inlet locates in a tropical 37 monsoon region with a distinct dry and wet season [10] coinciding with the Northeast 38 monsoon and Southwest monsoon, respectively [1]. From November to April, the dry 39 season is characterized by little rain, dryness, and moderate temperature (20-300C). The 40 wet season is from May to October and is characterized by high rainfall, humidity, and 41 high temperature (30-400c). The Tien Chau tidal inlet is filling with sediment [12]. The 42 deposition of sediment around the tidal inlet negatively influences the operation of boats 43 and ships in this region. It is suggested that the water depth at some regions in the ebb-44 tidal delta region is only about one meter during low-tide water. Only small boats of less 45

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than 50 horsepower can transport in these conditions, and local fishing vessels cannot
access the Tien Chau finishing port. This is a typical problem of tidal inlets located in the
central region of Vietnam [11].

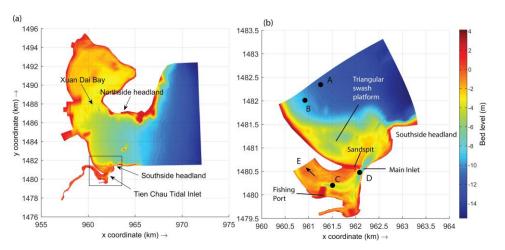


Figure 1. (a) Overall location of the study area, Tien Chau tidal inlet connects Tien Chau Estuary50and the South East Sea. (b) The zoom-in descriptions of the main entrance of the region with descriptions including the inlet, the Fishing Port, the updrift sand spit, the southside headland of the51Xuan Dai Bay. The detailed measured bathymetry of the study area (in 2019) and observation stations A, B, C, D, and E (in 2019) are included.53



Figure 2. Tien Chau tidal inlet captured by drone in 2019. It is observed that many houses are constructed at the downdrift barrier of the inlet. The sandspit at the updrift side moving toward the Southside.565758

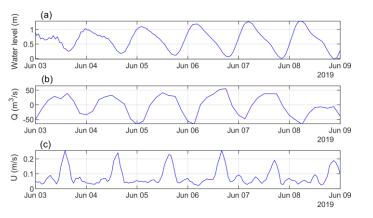
It is observed that the Tien Chau tidal inlet anchors next to a headland and is bor-59 dered on its opposite side by a sandspit (Figure 1 & 2). In this condition, most of the in-60 coming waves were blocked by the southside headlands (Figure 1 & 2). Moreover, the 61 river discharge during historical flooding events can be six times larger than that during 62 normal conditions [11]. Furthermore, it is observed that the human constructions, includ-63 ing the aquacultures and housing in the flood-tidal delta and in the downdrift barrier, 64 took the place of the flood-tidal delta. It is observed that the updrift sand spit of the tidal 65 inlet keeps fluctuating toward the ebb-tidal delta [11]. Understanding the hydrodynamics 66 and morphodynamics processes of monsoon tidal inlets that are affected by the river dis-67 charge, sheltered and constricted by human constructions and headlands is still in its in-68 fancy. Therefore, the main objectives of this study are to understand the primary morpho-69 logical and hydrodynamic processes of such inlet systems. 70

In order to obtain more insight into the physical processes of the study area, a field 73 measurement campaign was conducted. Important hydrodynamic and morphological 74 characteristics of the system were examined. In addition, a schematised model of the 75 study area was constructed in Delft3D. The model was calibrated and validated using the 76 measured data. The Tien Chau Estuary region's primary field measurement data sets are 77 the topography of the study area, the water level at the Tien Chau tidal inlet, and the flow 78 at the Tien Chau tidal inlet and the wave height at the Tien Chau coast. The current & 79 water level data were collected using Acoustic Wave and Current Profiler (AWAC). The 80 wave data was recorded using Wave droid. 81

Moreover, a schematised model of Tien Chau Estuary was constructed in Delft3D. 82 The model Delft3D was chosen as a numerical solver as it contains the necessary physical 83 processes, including the drying and wetting of large tidal flats, the interaction between 84 wave and current, and sediment transport [5]. Delft3D-FLOW has been commonly used 85 to simulate hydrodynamics and morphodynamics in shallow water environments, includ-86 ing tidal inlet systems [6]. 87

3. Results

Figure 3a illustrates the water levels at the Tien Chau fishing harbor (station C) from 89 2nd June and 9th June 2019. One high and one low tide per day can be seen in the time 90 series of the water level. The mean spring tidal range is around 1.5 m. Thus, the tidal 91 regime of the system can be classified as micro-tidal [4]. The form factor is 1.96, indicating that the Tien Chau estuary can be classified as mixed and mainly diurnal [2]. Moreover, 93 the falling period is larger than the rising period. Therefore, the Tien Chau tidal inlet ex-94 periences tidal asymmetry [3]. 95



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Figure 3. Typical forcing conditions at the Tien Chau tidal inlet, data collected between 2nd June 97 and 9th June 2019 including (a) Measured water levels at the fishing harbor (station C); (b) Hourly 98 discharge at the tidal inlet (station D); (c) depth-averaged streamwise velocity at the harbor (station 99 C). 100

Figure 3b shows the discharge time series measured by ADCP at station D, at the 101 main tidal inlet. The tidal prism (P) is estimated at approximately 2.4×10^6 m³ by P = 102 $\frac{1}{2}\int_0^T |Q(t)dt|$ [2]. P is the tidal prism, Q(t) is discharge at the inlet, and T is a tidal cycle. 103 Figure 3c illustrates the magnitude of the depth-averaged streamwise velocity at the fish-104 ing harbor. The maximum velocity is approximately 0.25 ms⁻¹. The peak currents at this 105 location typically occur during the low tide and are associated with the maximum dis-106 charge at the inlet entrance. 107

Figure 4. a, b sketch the major morphological characteristic features of a normal tidal 108 inlet located in a mesotidal/mixed energy region [2] and that of the Tien Chau tidal inlet, 109 respectively. In the Tien Chau tidal inlet, the downdrift barrier is absent, and the southeast 110

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headland provides a natural barrier blocking most of the incoming waves from the north-111 east direction. As a result, in the ebb-tidal delta of the Tien Chau tidal inlet (Figure 4b), 112 the left and right swash bars, which are usually formed in a usual tidal inlet (Figure 4a), 113 are merged into a single large triangular swash bar platform (see Figure 1b and 4). A ter-114 minal lobe is located at the outer edge of the triangular swash bar. In the flood-tidal delta, 115 different tidal flats also tend to migrate and merge into a large tidal flat region (Fig 4b). 116 This tidal flat region is directly attached to the sand spit and located next to the main 117 channel of the inlet. 118

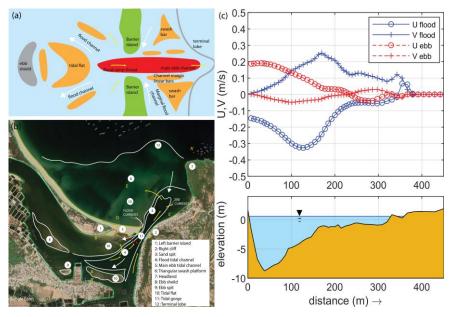


Figure 4. The major morphological features of (a) a regular tidal inlet located in a mesotidal/mixed120energy (Bosboom and Stive, 2021); (b) the Tien Chau inlet: an asymmetry, sheltered and squeezed121located in a monsoon region with micro-tidal/moderate energy; (c) Profiles of depth-averaged122streamwise velocities (circles) & transverse velocities (plusses) extracted from the numerical model123during the flood period (blue) and ebb period (red) near the throat of the inlet (section C-C').124

Moreover, as described previously, it is also observed that the construction of houses, 125 fish farms, and fishing port blocks the development of the downdrift barrier and limits 126 the accommodation space in the flood-tidal delta. In these conditions, the spatial re-127 striction disturbs the dynamic equilibrium of the tidal inlet system and prevents the inlet 128 from "fluctuating" toward its new equilibrium stage under the combined influences of 129 sea-level is rises and annual river floods. The sand spit and the accommodation space of 130 the system tend to be pushed toward the outer delta. Assuming that there is no change in 131 the longshore drift, then the amount of the sediment accumulated in the outer delta will 132 increase continuously, and the tidal inlet is continuously filling with sediment. In this 133 sense, the term squeeze, which has been widely used for coastal regions [7] and ecological 134 systems [8], can also be adopted for the Tien Chau tidal inlet. 135

Figure 4c shows the numerical results of the distribution of the depth-averaged 136 streamwise and transverse velocity at the crosssection C-C' near the throat of the inlet (see 137 Figure 4b). It is suggested that the water depth in the main channel increases to 8.5 m. 138 During the flood period, the depth-averaged streamwise velocity reaches its maximum 139 value of around 0.32 ms⁻¹ at the middle of the transition slope between the shallower tidal 140 flat and the main channel (x = 120 m). The depth-averaged transverse velocity is positive, 141 and its maximum value is about 0.3 ms⁻¹. This means that the direction of the flood cur-142 rents is also toward the left bank. It is observed that a single main channel is flanked by 143 the flat area on the flood tidal delta and the triangular swash region on the ebb-tidal delta. 144

In this context, the hydrodynamics of the flood-tidal delta is similar to that of a compound channel (see lower panel of Figure 4c) consisting of a floodplain, transition slopes, 146

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and a main open channel [9]. Thus, from a fluid mechanics perspective, the squeeze phe-147 nomenon of the Tien Chau tidal inlet may be approached as similar to the squeeze of a 148floodplain channel. The width of the floodplain region was reduced due to the construction of fish farms [9]. 150

5. Conclusions

This study investigated the forcing conditions, morphological features, and hydro-152 dynamic characteristics of an asymmetry tidal inlet located in a monsoon region, sheltered 153 by headlands, and affected by human constructions. Morphological features of the inlet 154 were identified and studied using measured data in combination with satellite imagery. 155 A numerical model was built to determine the depth-averaged streamwise and transverse 156 velocity profile at the inlet. It is suggested that the north and southeast headlands of Xuan 157 Dai Bay offer a natural barrier blocking most of the southeast and north coming waves. A 158 wave-built sand spit elongated from the updrift sand barrier island can be clearly ob-159 served. Moreover, it is observed that the human constructions block the development of 160 the downdrift sand barrier island and the accommodation space in the flood-tidal delta. 161 In this context, the term squeeze was adopted for the Tien Chau tidal inlet. This phenom-162 enon makes the morphological features of the Tien Chau tidal inlet different from that of 163 a regular inlet. Small tidal flats tend to merge to a large one and are more dominant 164 around the inlet and the ebb-tidal delta. The triangular swash bar in the ebb-tidal delta 165 and the tidal flats in the flood-tidal delta are located next to the main inlet, forming a 166 compound channel. The sand spit tends to fluctuate toward the outer delta, and more 167 sediment tends to be deposited in the ebb-tidal delta. Now the generic understanding of 168the hydrodynamic and morphological features of the Tien Chau tidal inlet has been clari-169 fied. Future work will further prove the squeeze hypothesis adopted for the Tien Chau 170 tidal inlet by applying different scenarios of the numerical simulations. 171

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