Modelling External Risk to Agricultural Development Projects using A Reconnaissance Drought (RDI) and Standardised Precipitation Indices (SPI)

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Abstract: Drought is defined as a recurrent feature of climate that involve deficiency of precipitation over an extended period resulting in water shortages that have adverse effects on economy, environment and humans and their health. The probability of drought impacts varies widely depending upon drought duration, frequency and severity, baseline population vulnerability, existing health and sanitation infrastructure, and available resources with which to mitigate impacts as they occur. This natural climate feature adversely leads to poor crop yields. The current study aims at charactering drought for proactive management and mitigation strategies to be applied. The study used standardised precipitation (SPI) and reconnaissance drought (RDI) indices to model the behaviour of drought in the study area. The key findings revealed the study area under frequent, severe to extreme drought events. The study recommends that water users and all other relevant stakeholders adopt mitigation strategies such as use of water sparingly, use of drought-resistant crop cultivars and others.

Keywords: Project risk; risk modelling; RDI; SPI

1. Introduction

Drought is described as a period in time whereby conditions are drier than normal and there is less rainfall than usual for a long period of time; this could be for several weeks, months and years which lead to water-related problems (Praveen, Ramachandran, Jaganathan, Krishnaveni, & Palanivelu, 2016). Needless to say, there is no specific definition of drought; however it can be classified into four main types. In this study, drought is described as a severe condition which is characterized by a dry spell which has negative impacts on the economy, society and environment. The four main types of drought are: meteorological drought, hydrological drought, agricultural drought and socio-economic drought. (i) Meteorological drought is referred to as the degree of dryness compared to normal conditions and this is uneven over time and particular regions (Olivares Campos, 2016), (ii) hydrological drought refers to insufficient rainfall which leads to serious reduction in run-off streamflow, inflow into storage reservoirs and revitalizing of groundwater (Orlando Olivares & Zingaretti, 2018), (iii) Agricultural drought refers to a deficit of soil moisture to sustain plants and livestock thereby causing redundant growth and reduced produce (Tefera, Ayoade, & Bello, 2019) and Socio-economic drought refers to a period in which human activities are poorly affected by reduction in water availability and precipitation. (Praveen et al., 2016). Much as all these types of drought have a connection, one leads to another and sometimes overlaps; so, agricultural drought is regarded as of prime importance and a roadmap in this study. The diagram below illustrates the interrelationship of the types of drought.
1.1. Key Driving Forces Behind Drought Events

Of all the hemispheres (north, south of the equator; east and west of the Greenwich Meridian), Africa is the only continent that lies within them all and is affected by numerous climatic conditions. (Roman-Cuesta, 2007). Consequently, this continent receives conventional rainfall. (Choi, An Prof., Yeh, & Yu, 2013). This also gives rise to different climatic conditions such as El Nino. El Nino is a recurring climate pattern which is caused by changes in temperatures of water by warming the eastern Pacific, and therefore affecting the global climate. (Ali & Ali, 2011). This natural occurrence is responsible for exacerbating drought events in the southern hemisphere. (Ali & Ali, 2011). In accordance to Keil, Zeller, Wida, Sanim, & Birner (2008), drought is a resultant of the shifting, changing weather patterns. On the other hand, it is also a belief that droughts chiefly transpire because of natural occurrences due to the earth and atmospheric systems. However, Granzow-de la Cerda, Lloret, Ruiz & Vandermeer (2012) are of the belief that no one knows for sure why droughts occur, despite the fact that many scientists believe that there is a relationship between drought occurrence and El Nino events. Enhanced understanding of the relationships between droughts and repeated changes in high and low pressures from one side of the Pacific to the other linked with La Nina allow scientists to formulate improved predictions of this cataclysmic drought hazard. (Ryu, Svoboda, Lenters, Tadesse, Knutson 2010). Both El Nino and La Nina form the El Nino-Southern Oscillation (ENSO) cycle. ENSO is a recurring climate pattern whereby temperature fluctuate between the ocean and atmosphere in the east-central Equatorial Pacific, approximately between the International Date Line and 120 degrees West (Olivares Campos, 2016). Nevertheless, these two events happen every 2-7 years with El Nino events occurring more often than those of La Nina (Tefera et al., 2019). There has been a rising trend in global weather disasters since 1980, and with extreme climatic events such as droughts. (Orlando Olivares & Zingaretti, 2018). Comprehending phenomena such as ENSO is pivotal because of the possibility that it could cause an enormous loss of property, destruction to the environment and the loss of human life. Forecasts by ENSO, tracking and monitoring play a significant role to insurers, government authorities and other pertinent stakeholders for drought management for proactive planning against unfavourable impacts such as acute climatic changes like drought events (Praveen et al., 2016).
1.2. Solar Activity (Sunspots Number)

Drought could also be associated with sunspots which commonly last for a period of approximately 11 years. (Minckley, Roulston, & Williams, 2013). Sunspots are described as cool surface areas on the sun that are visible in pairs and are darker in comparison with other parts of the sun. These spots have a strong magnetic field and rotate like a giant hurricane. (Mèthy, Damesin, & Rambal, 1996). These authors also affirm that several authors have acknowledged that sunspots have impacts on temperature, precipitation, length of growing seasons, air circulation, atmospheric pressure, high altitude, wind speed and other natural phenomena around the world. Moreover, Xiao & Zhuang (2007) construct solar activity as a main cause of cyclic deviations of the global climate through triggering of the evaporation processes. Scientists have also shown that sunspot numbers and drought events are correlated. For instance, during solar activity- drought events take place at solar maximum (Solheim, Stordahl, & Humlum, 2011). This occurrence is related to climatic conditions where temperatures become high during solar maximum (Minnis, 1958). Solar energy is the principal energy source as well as control on evaporation; therefore, distributions of insulation and evaporation are strongly linked. (Siingh et al., 2011). The energy from the sun is the central source of energy present for heating the surface of the planet earth. This energy supplied by the sun is an outcome of its activity and it differs with time. The major cycle of the sun is eleven years. The major cause of drought events is believed to be the solar activity (Ghormar, 2014). The coefficient of correlation between insolation and evaporation ranges between 0.820 and 0.948 and values of the calculated solar radiation are used in the computation of the Potential Evapotranspiration (PET) in Penman equation (Abarca del Rio, Gambis, Salstein, Nelson, & Dai, 2003). The solar radiation that falls on the earth’s surface depends on the distance it travels to the object and the angle at which rays hit an area or object. (Méthy et al., 1996). The universal law for the intensity of radiation, distinctively the sine law of sunlight states that the sunlight always strikes the high latitude obliquely, so it spreads out more and is less intense. (Minckley et al., 2013).

1.3. Drought Impacts

Drought has several impacts, and these include mass starvation, famine and a pause or sometimes an end to economic activity particularly in areas where rain fed agriculture is the mainstay of the rural economy (4, n.d.). It is generally known that drought is the chief cause of forced human migration and environmental refugees, deadly conflicts over the use of diminishing natural resources, food insecurity and starvation, a damage to significant habitations and as well as loss of biological diversity; volatility of socio-economic conditions, poverty and unpredictable climatic conditions through reduced carbon sequestration possibility (Roman-Cuesta, 2007). Drought and desertification impacts are among the pricey incidents or occurrences in Africa, for instance, the prevalent destitution as well as the unstable economy of many African countries which in actuality depend on climate-sensitive segments such as rain-fed agriculture. Drought and desertification impacts are among the pricey incidents or occurrences in Africa, for instance, the prevalent destitution as well as the unstable economy of many African countries which in actuality depend on climate-sensitive segments such as rain-fed agriculture. All plants and animal life present in a particular region which are not resistant to drought are most likely to go extinct. (Nagamuthu & Rajendram, 2015). The collective results of drought and bush burning (during dry seasons) have made the plants to go extinct and animals to drift into safer places. Drought, land degradation and desertification have had grave impacts on the richness and variety of fauna and flora (Francisco, 2013). Moreover, plants biodiversity will alter with time, unpleasant species will dominate, and total biomass production will dwindle (Khan & Gomes, 2019). Plants and animals are dependent on water, like people. Drought can minimize their food supplies and damage their habitats. Occasionally, this damage lasts for only a limited period of time, and other times it is irrevocable. Drought can also affect people’s health and safety. For example, drought impacts on society include anxiety or depression about economic losses, conflicts due to water shortages, reduction of income, fewer recreational activities, and increase of heat stroke incidents and sometimes loss of human lives. Moreover, drought conditions can also grant a considerable increase in wildfire risk. This is due to
withering of plants and trees from lack of precipitation, scourge insect infestations, and diseases, all of which are associated with drought. (Prokurat, 2015). Lengthy periods of drought can cause more wildfires and more powerful wildfires, which impinge on the economy, the environment as well as the society in a number of ways like destroying neighbourhoods, crops and habitats (Do Amaral, Cunha, Marchezini, Lindoso, Saito, & Dos Santos Alvala, 2019). Moreover, drought not only always offers similar instant and remarkable visuals related with occurrences such as hurricanes and tornadoes, but it still has a huge price tag. In point of fact, droughts are the second in rank types of phenomena that are associated with billion-dollar weather disasters for the past three decades (Nagamuthu & Rajendram, 2015). With staggering yearly losses close to $9 billion annually, drought is a severe hazard with socioeconomic risks for most African countries. (Siingh et al., 2011). These pricey drought impacts come in different forms. For instance, the economic impacts of drought include farmers who lose money because drought destroyed their crops or worse ranchers who may have to spend more money for animal feeds and irrigation of their crops. Economic impacts can either be direct or indirect. Directly, it could be a decrease in dairy production and indirectly, it could be observed through increases in the price of the cheese (Francisco, 2013).

2. Materials and Methods

The monthly rainfall dataset of this study was obtained from National Aeronautics and Space Administration (NASA) data portal. This dataset was used as the only input parameter for Standardized Precipitation Index computation. Standardized Precipitation Index (SPI) is plainly described as a normalised index that signifies a likelihood of a rainfall occurrence of an observed rainfall amount in comparison with the rainfall climatology at a particular geographical location over a long-term reference period (Singh et al., 2011). Furthermore, Yusof, Hui-Mean, Suhaila, Yusop, & Ching-Yee (2014) affirm that SPI is a probability index that offers an enhanced demonstration of both abnormal wetness and dryness than any Palmer indices such as Palmer Drought Severity Index (PDSI). The value of this index is that it can be computed for different time scales, for that reason, issuing early warming of drought and its severity (Gaas, 2018). This index is appropriate for risk management purposes (Verma, Verma, Yadu, & Murmu, 2016). Moreover, this index is advantageous in that precipitation is the only parameter in its computation therefore making it less complex. Conversely, the weakness of this index is that it can only compute the precipitation deficit; values found on initial data may alter, and values vary as the period of record grows (Jordan, 2017). The table 1 below illustrates the values of SPI categorisation.

<table>
<thead>
<tr>
<th>SPI Value</th>
<th>Category</th>
<th>Probability %</th>
</tr>
</thead>
<tbody>
<tr>
<td>≥2.0</td>
<td>Extremely wet</td>
<td>2.3</td>
</tr>
<tr>
<td>1.5 to 1.99</td>
<td>Very wet</td>
<td>4.4</td>
</tr>
<tr>
<td>1 to 1.49</td>
<td>Moderately wet</td>
<td>9.2</td>
</tr>
<tr>
<td>-0.99 to 0.99</td>
<td>Near normal</td>
<td>34.1</td>
</tr>
<tr>
<td>-1.0 to -1.49</td>
<td>Moderately dry</td>
<td>9.2</td>
</tr>
<tr>
<td>-1.5 to -1.99</td>
<td>Severely dry</td>
<td>4.4</td>
</tr>
<tr>
<td>≤-2.0</td>
<td>Extremely dry</td>
<td>2.3</td>
</tr>
</tbody>
</table>

Source: Hlalele, 2016

2.1. Methodology

The SPI calculations are founded on the fact that precipitation increases over a fixed time scale of interest, for instance; SPI-3, SPI-6, SPI-9, SPI-12, SPI-24 and SPI-48, so from that a series is integrated in a gamma probability distribution which is apt for this climatological precipitation time series. (Yusof et al., 2014). The gamma distribution is described by the following density function.

\[ g(x) = \frac{1}{\beta^\alpha \Gamma(\alpha)} x^{\alpha-1} e^{-x/\beta} \quad \text{for } x > 0 \]  

(1)
where $\alpha$ and $\beta$ are estimated for each station as well as for each month of the year.

\[
\alpha = \frac{1}{4A} \left( 1 + \sqrt{1 + \frac{4A}{3}} \right)
\]

\[
\beta = \frac{\overline{x}}{\alpha}
\]  

where $A = \ln(x) - \frac{\ln(x)}{n}$, and $n$ = number of observations

After these parameters have been estimated then their resulting values are used to calculate cumulative probability as

\[
G(x) = \frac{1}{\beta^\alpha \Gamma(\alpha)} \int_0^x x^{\alpha-1} e^{-x/\beta} dx
\]  

In cases where $t=x/\beta$ then an incomplete gamma function becomes

\[
G(x) = \frac{1}{\Gamma(\alpha)} \int_0^x t^{\alpha-1} e^{-t} dt
\]

Since gamma function is undefined at $x=0$ then the cumulative probability is calculated from the following equation (Rahmat, Jayasuriya, & Bhuiyan, 2012):

\[
H(x) = q + (1 - q) G(x),
\]

where $q$ is the probability of a zero and $G(x)$ the cumulative probability of the incomplete gamma function. If $m$ is the number of zeros in a precipitation time series, then $q$ can be estimated by $m/n$. The cumulative probability is then transformed to the standard normal random variable $z$ with mean zero and variance one, which is the value of the SPI (Yusof et al., 2014). In the event where it is standardized, the potency of the irregularity is categorised as illustrated in Table xx, where the table also demonstrates the corresponding probabilities.

2.2. Data Quality Control

Statistically, homogeneity tests are carried out to scrutinize statistical properties of a certain dataset. In essence, it thoroughly looks at the location stability and variations which are local within the time series over time (Abraham & Yatawara, 1998). The author also confirms that this occurrence is the same as testing statistical distribution, for that reason identifying if there are any changes in the distribution. The test is also carried out to evade false or unauthentic results from the data sets. (Hosseinzadeh Talaee, Kouchakzadeh, & Shifter Some’e, 2014). A non-parametric Pettitt’s test was used. Outliers and missing were identified and substituted by Expectation Maximum algorithm (EIM) with the help of SPSS software. EM is described as a statistical algorithm appropriate to be used when there are missing or hidden values in the data sets (Lobato & Velasco, 2004). Tan & Yilmaz (2002) construct that EM is a well-liked too used in statistical estimation problems that involve data which is incomplete. Likewise, Technology & Bay (2001) define EM as an algorithm that allows parameter estimation in probabilistic models with data which is not complete.

Before carrying out any data analysis, it is essential to assess the apparent proof of patterns and trends in the climate data (Kliewer & Mertins, 1997). Non-parametric Mann-Kendall test is used in the study to assess if any trends existed. This test is universally used to identify monotonic trends in series of environmental, climate and hydrological dataset. The null hypothesis (H0) means that data came from a population with autonomous realisations are identically distributed. The alternative hypothesis (Ha), means that data follows a monotonic trend. The Mann-Kendall statistic indicates how strong and weak two variables are associated and show correlation direction.(Kliewer & Mertins, 1997). One of the advantages of this statistic is that the data does not necessarily have to
follow any specific probability distribution. The test was conducted simultaneously with a non-parametric Pettitt’s test to gauge data homogeneity and descriptive statistics.

2.3. Parameters used to Characterise Drought

The temporal characteristics are those features of a hazard associated with time and they are commonly linked with questions such as the following: When does the hazard occur? What is the frequency of the occurrence? What is the duration of the hazard? How fast do they hit and how conventional are they? (Andreadis, 2005: Van Niekerk, 2011). Drought intensity is described in numerous ways by different academics; nevertheless, in accordance to Pope et al (2013) intensity is a degree of insufficient rainfall. The authors further explain that, intensity can be defined as a result of duration as well as intensity. Abdulmaleket al. (2013) affirms that drought intensity gauges how far rainfall is below the average precipitation of the region. Understanding intensity can be used as a way of ascertaining the feasible impact of a hazard on communities. Understanding intensity can be used as a way of ascertaining the feasible impact of a hazard on communities as well as the levels of risk at which elements are exposed to (Van Niekerk, 2011). This aspect of drought is conveyed in several parameters such as the Standard Total Accumulative Dry Spell (STCD), Average Dry Spell Index (ADSI), Longest Multi year Drought (LMYD) and Largest Single Year Drought (LSYD) (Abdulmaleket et al., 2013). STCD signify the total cumulative drought index used. One more parameter used to quantify the same aspect is ADSI. LMYD and LSYD are other parameters obtained from drought indices such as SPI whose high values have negative outcomes on every facet of the environment, including socio-economic situation of communities. ADSI values offer valuable knowledge on the region’s characteristics essential for arrangement of water resources as well as irrigation projects. Areas with low values need special attention. Likewise, The ADSI values are of great significance to decision makers for the planning of agricultural projects of the affected areas for future. The LSYD also is significant to take into consideration in crop cultivation given that crops barely survive its high values. So, these four parameters are defined by the equations below:

\[
\text{ADSI} = \frac{\text{STCD}}{N}
\]

where \(N\) = total number of years of the series

(6)

\[
\text{LMYD} = \text{Maximum of any successive years}
\]

(7)

\[
\text{LSYD} = \text{Maximum drought index value of the single year}
\]

(8)

2.3.1. Drought Frequency and Duration Analysis

For a long-term planning to be effective in water projects such as irrigation and dam sizing purposes, there has to be an analysis of both dry and wet spells from a climatic and hydro-meteorological standpoint. (Abdulmaleket al., 2013). In drought analysis a period declared as dry when SPI< 0. These are some of the parameters used to calculate the drought duration: Longest Dry Spell Duration (LDSD), Drought Tendency (DT) and Average Dry Spell Duration (ADSD). LDSD is described as the highest of any consecutive dry spells that occurred on one occasion through the study record \(N\). High values of both LDSD and ADSD it means that water resources planning must be considered in that particular region. Nevertheless, DT is the ratio which involves the total number of dry spell cases to the wet spell cases. This parameter measures the predisposition of the study area which suffers from the dry spells, thus it measures the frequency of a hazard under consideration. These are defined by the following equations:

\[
\text{LDSD} = \sum_{i=1}^{N} D_i
\]

(9)

\[
\text{DT} = \frac{\sum D}{\sum W}
\]

(10)

\[
\text{ADSD} = \frac{\sum D}{N}
\]

(11)
If successive dry spell cases (D) are followed by a wet spell, like D, D, D, W, D then \( \sum D = 4 \), and \( N = 2 \) since an interrupted sequence of several cases of (D) constitute only one dry spell event.

A frequency analysis offers an early warning system. Disaster managers and appropriate stakeholders have the ability to foresee when the next incident will take place. (Sobrino et al. 2011). Hydro-climatic hazards such as drought have a propensity of following a seasonal pattern. When the number of times and the length of a hazard such as drought are known, such knowledge helps in planning accordingly (Hlalele, 2016). In a frequency analysis, approximation of the probability of an incidence of future occurrences is established on a primary base for risk management. (Yuliang et al. 2014). Again, it is used to foresee how frequently a hazard event happens over space and time (Des Jardins, 2012).

3. Results and Discussion

Standardised precipitation drought index is a world-renowned drought index used to quantify and characterise drought events. In this study drought has been analysed using SPI which uses only precipitation as an input parameter. RDI was only used which is an indicator that uses both precipitation and potential evapotranspiration (PET) from temperature-based equations. The two indices were used in order to complement each other. RDI is a water balance index used to quantify and characterise drought events. Figures 2 to 6 show the plots of seasonal and annual SPI’s respectively. Figure 2 shows a graph of Summer SPI-3 which has detected 2015/16 hydrological year as the driest year over the study record. Figures 3 and 4 indicated the 1992/3 as their driest years while 2013/14 was the driest in Spring. SPI/RDI-12 plot on figure 6 also showed 2015/16 as the driest year. The 2015/16, was the driest year in South Africa where the whole country was declared under drought disaster. It can therefore be seen that the summer drought events influence the annual drought in the study area. For all graphs SPI and RDI correlated strongly with a non-parametric Spearman’s coefficient of 0.99. However, on all temporal scales, SPI has shown to do slightly better than RDI where it detected drought severe drought events. Table 2 below shows hydrological years detected by all the four common selected seasons and annual SPI. At extreme levels Summer drought events are same as those observed by annual SPI. The 2015/16 was the year in which drought was a disaster in the study area which is largely accounted for by summer droughts. Several drought parameters as shown in table 3 were identified to assess the levels of frequency, average duration and intensity. Drought tendency (DT) measures the ratio of dry/drought spells (D) to wet spells (W). Annual, Spring and Summer SPI’s have the highest DT above 1, which implies that the study area experiences frequent dry spells than any other season. The average drought/dry spell duration for all the selected time scales, seasonal and annual is approximately 2 months and two years respectively which are in the severe drought category as depicted in table 4.
Figure 3. Bloemfontein Autumn SPI/RDI-3 plot (1977-2017).

Figure 4. Bloemfontein Winter SPI/RDI-3 plot (1977-2017).

Figure 5. Bloemfontein Spring SPI/RDI-3 plot (1977-2017).
Figure 6. Bloemfontein Annual SPI/RDI-3 plot (1977-2017).

Table 2. Drought Years From 1977-2017 Per SPI/RDI Drought Category.

<table>
<thead>
<tr>
<th>Time scale</th>
<th>Moderate</th>
<th>Severe</th>
<th>Extreme</th>
</tr>
</thead>
<tbody>
<tr>
<td>Spring</td>
<td>1986/87, 2010/11</td>
<td></td>
<td></td>
</tr>
</tbody>
</table>

Table 3. Drought Parameters.

<table>
<thead>
<tr>
<th>Time scale</th>
<th>Season</th>
<th>D</th>
<th>W</th>
<th>DT</th>
<th>N</th>
<th>ADSD</th>
<th>STCD</th>
<th>ADSI</th>
</tr>
</thead>
<tbody>
<tr>
<td>SPI/RDI-3</td>
<td>Summer</td>
<td>25.00</td>
<td>15.00</td>
<td>1.67</td>
<td>10.00</td>
<td>2.50</td>
<td>-15.57</td>
<td>-1.56</td>
</tr>
<tr>
<td></td>
<td>Winter</td>
<td>17.00</td>
<td>23.00</td>
<td>0.74</td>
<td>9.00</td>
<td>1.89</td>
<td>-14.14</td>
<td>-1.57</td>
</tr>
<tr>
<td></td>
<td>Spring</td>
<td>24.00</td>
<td>16.00</td>
<td>1.50</td>
<td>10.00</td>
<td>2.40</td>
<td>-15.77</td>
<td>-1.58</td>
</tr>
<tr>
<td></td>
<td>Autumn</td>
<td>18.00</td>
<td>22.00</td>
<td>0.82</td>
<td>9.00</td>
<td>2.00</td>
<td>-16.54</td>
<td>-1.84</td>
</tr>
<tr>
<td>SPI/RDI-12</td>
<td>Annual</td>
<td>22.00</td>
<td>18.00</td>
<td>1.22</td>
<td>10.00</td>
<td>2.20</td>
<td>-16.70</td>
<td>-1.67</td>
</tr>
</tbody>
</table>

Table 4. SPI/RDI Values Categories.

<table>
<thead>
<tr>
<th>SPI Value</th>
<th>Category</th>
<th>Probability %</th>
</tr>
</thead>
<tbody>
<tr>
<td>&gt;2.0</td>
<td>Extremely wet</td>
<td>2.3</td>
</tr>
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<td>1.5 to 1.99</td>
<td>Very wet</td>
<td>4.4</td>
</tr>
<tr>
<td>1 to 1.49</td>
<td>Moderately wet</td>
<td>9.2</td>
</tr>
<tr>
<td>-0.99 to 0.99</td>
<td>Near normal</td>
<td>34.1</td>
</tr>
<tr>
<td>-1.0 to -1.49</td>
<td>Moderately dry</td>
<td>9.2</td>
</tr>
<tr>
<td>-1.5 to -1.99</td>
<td>Severely dry</td>
<td>4.4</td>
</tr>
<tr>
<td>&lt;-2.0</td>
<td>Extremely dry</td>
<td>2.3</td>
</tr>
</tbody>
</table>

4. Conclusion and Recommendations

In conclusion, Drought is defined as a recurrent feature of climate that involve deficiency of precipitation over an extended period resulting in water shortages that have adverse effects on economy, environment and humans and their health. The probability of drought impacts varies widely depending upon drought duration, frequency and severity, baseline population vulnerability, existing health and sanitation infrastructure, and available resources with which to mitigate impacts as they occur. This natural climate feature adversely leads to poor crop yields. The current study aims
at characterizing drought for proactive management and mitigation strategies to be applied. The study used standardised precipitation (SPI) and reconnaissance drought (RDI) indices to model the behaviour of drought in the study area. The key findings revealed the study area under frequent, severe to extreme drought events. The study recommends that water users and all other relevant stakeholders adopt mitigation strategies such as use of water sparingly, use of drought-resistant crop cultivars and others.

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