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Soil-water dynamics in flood irrigated orange orchard in central India: Integrated approach of sap flow measurements and HYDRUS 1D model

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# Outline



Genesis of the Study











# **Genesis of the Study**

- The N-P CZO is intensively managed watershed
- 60% land utilised for agri-horticulture
- Extensive use of GW for irrigation of orange orchards
- Watershed is under overexploited condition (GW stage development >100 %)





### HYDRUS 1D model

- Assessment of water loss due to evaporation and deep drainage in the present scenario
- Understanding pattern of root water uptake
- Sensitivity analysis of the parameters
- Optimization of irrigation schedule

# **HYDRUS-1D Model Equations**

Richards Equation for water flow and root water uptake in variably saturated soil:

$$\frac{\partial \theta}{\partial t} = \frac{\partial}{\partial z} \left[ k \left( \frac{\partial h}{\partial z} + 1 \right) \right] - \mathbf{S}(z, t)$$

The sink term can be defined as:

$$\mathbf{S}(\mathbf{h}) = \boldsymbol{\alpha}(\mathbf{h}) \frac{\mathbf{b}'(z)}{\int_0^{\mathbf{L}_{\mathbf{r}}} \mathbf{b}'(z) dz} \mathbf{T}_{\mathbf{p}}$$

Relationship between  $\theta$  and h, and K and  $\theta$ 

(van Genuchten-Mualem 1980):

$$\boldsymbol{\theta}(\boldsymbol{h}) = \begin{cases} \boldsymbol{\theta}_r + \frac{\boldsymbol{\theta}_s - \boldsymbol{\theta}_r}{[1 + |\boldsymbol{\alpha}\boldsymbol{h}|^n]^m} & \boldsymbol{h} < 0\\ \boldsymbol{\theta}_s & \boldsymbol{h} \ge \mathbf{0} \end{cases}$$

$$K(h) = K_s \left(\frac{\theta - \theta_r}{\theta_s - \theta_r}\right)^l \left\{ 1 - \left\{ 1 - \left(\frac{\theta - \theta_r}{\theta_s - \theta_r}\right)^{\frac{1}{m}} \right\}^m \right\}^m$$

 $\theta$  = Volumetric soil-water content [L<sup>3</sup>L<sup>-3</sup>] *h* = Soil-water pressure head [L] k = Unsaturated soil hydraulic conductivity [LT<sup>-1</sup>] z = The spatial coordinate (positive upward) [L] **t** = **Time [T]**  $S = Sink term [L^3 L^{-3} T^{-1}]$  $\alpha(h)$  = Dimensionless function ( $0 \le \alpha \le 1$ ) L<sub>r</sub> = Rooting depth [L]  $\mathbf{b}'(\mathbf{z}) = \mathbf{Root}$  distribution function  $[\mathbf{L}^{-1}]$  $T_p$  = Potential transpiration rate [LT<sup>-1</sup>]  $\theta_r$  = residual water content  $\theta_s$  = saturated water content  $\alpha$  = inverse of the air-entry value (or bubbling pressure) *n* = pore-size distribution index  $K_s$  = saturated hydraulic conductivity l = pore-connectivity parameter, and m = 1 - 1/n

### **Geometry Information**

- Number of Soil Material: 01
- Depth of Soil Profile: 3 m

### **Model Selection**

- Soil Hydraulic Model: van Genuchten-Mualem
- Root Water Uptake Model: Feddes

### **Simulation Period: 61 days**

- 4 wet periods (04 days each)
- 3 dry periods (15 days each)



### **Feddes Curve for Orange Tree**



### **Estimation of Soil Hydraulic properties**

- Soil Hydraulic properties: Estimated in HYDRUS 1D by using soil texture data\*
- The soil texture of the study site (clay: 60%, silt: 25% and sand 15%)\*

\* Rosetta Dynamically Linked Library (DLL)
\* Soil sample was collected from 15 cm below ground and Texture Analysis was done in Lab

Soil Hydraulic Parameter	Units	Value
Saturated hydraulic conductivity ( $K_s$ )	cm d⁻¹	21.12
Saturated soil water content ( $\boldsymbol{\theta}_s$ )	-	0.4957
Residual soil water content ( $oldsymbol{ heta}_r$ )	-	0.0992
The inverse of the air-entry value	d <sup>-1</sup>	0 0101
(or bubbling pressure) ( $lpha$ )	u	0.0191
Pore-size distribution index ( <i>n</i> )	-	1.224
Tortuosity parameter in the conductivity function ( <i>l</i> )	-	0.5 7

### **Boundary Conditions**

- Upper Boundary Condition: Variable pressure head/flux
- Lower Boundary Condition: Free drainage

### **Input parameters**

- Top flux: (Irrigated water +Rainfall)-Evaporation\*
- Daily Potential Transpiration\*
- Root Distribution- Assumed to be (90 cm)<sup>†</sup> and extends horizontally up to 2 m



### **Calculation of Irrigated water per day (Flood Irrigation)**

Irrigated on 1 Acre of orchard (198 trees) for 12 h in a day:

$$1 Acer = 4046.87 m^2$$

**Discharge rate of Bore well** =  $3.2 \frac{l}{s}$ 

$$3.2 \frac{l}{s} \times 60 \times 60 = 11520 \frac{l}{h}$$

Water is applied for 12 h in a day:

$$11520 \frac{l}{h} \times 12 h = 138240 \frac{l}{d} = 138.24 \frac{m^3}{d}$$
Applied water on unit area  $= \frac{138.24 \frac{m^3}{d}}{4046.87 m^2} = 0.034156 \frac{m}{d} = 3.42 \frac{cm}{d}$ 



### Scaling down sap flow data

- The circumference (S<sub>T</sub>) of the orange tree was 0.52 m
- The sapwood area (S<sub>A</sub>) was calculated as 0.30 m<sup>2</sup> (Granier A., 1987)

 $S_A = -0.0039 + 0.59 S_T$ 

### Scaled Sap Velocity (V)

 $\mathbf{V} = (\mathbf{S}_{\mathrm{A}} \times \mathbf{V}_{\mathrm{sap}})/\pi \ \mathbf{r}^2$ 

r = Radius of root spread

 $V_{sap} = Measured sap velocity$ 



### **Experimental Set-up**



Thermal Dissipation Probe (TDP) sensors Make: Dynamax Inc., U.S.A.

CR1000X Measurement and Control Datalogger, Make: Campbell Scientific, Inc., U.S.A



5 Year old orange tree (Young) Tree Height: 2.7 m Circumference: 25 cm 15 Year old orange tree (Mature)

Tree Height: 3.4 m Circumference: 52 cm







Automatic Weather Station Make: Rainwise Inc. U.S.A. 11

# **Estimation of Potential Evapotranspiration Rate**

#### **Parameter** Source **Penman-Monteith equation Meteorological Parameters** $ET_{o} = \frac{1}{\lambda} \left[ \frac{\Delta(R_{n} - G)}{\Delta + \gamma (1 + r_{c}/r_{a})} + \frac{\rho c_{p}(e_{a} - e_{d})/r_{a}}{\Delta + \gamma (1 + r_{c}/r_{a})} \right]$ (Temperature, Relative Humidity, Wind Speed, Weather **Station** Solar Radiation) **ET**<sub>o</sub> = **Potential evapotranspiration rate** Soil Heat Flux $\lambda$ = Latent heat of vaporization **NASA Satellite Cloud Fraction** Data $\Delta$ = Slope of the vapor pressure curve **Net Heat Flux** $\mathbf{R}_{\mathbf{n}} = \mathbf{Net} \mathbf{radiation} \mathbf{at} \mathbf{surface}$ **G** = Soil heat flux $\gamma$ = Psychrometric constant **r**<sub>c</sub> = **Crop canopy resistance** 1.2 **r**<sub>a</sub> = Aerodynamic resistance DOY DOY $\rho$ = Atmospheric density 756 **c**<sub>n</sub>= Specific heat of moist air а (2.600 E (cm d<sup>-1</sup>) 550 $e_a$ = Saturation vapor pressure at T 450 400 $e_d$ = Actual vapor pressure 350 300 30 DOY DOY

# **Partitioning and diurnal variation-PET**

#### Partitioning of Potential Transpiration and Evaporation

$$T_{p} = ET_{p} (1 - e^{-k \text{ LAI}})$$
$$E_{p} = ET_{p} e^{-k \text{ LAI}}$$



 $ET_p$  = Potential evapotranspiration,  $E_p$  = Pot. Evaporation,  $T_p$  = Pot. Transpiration

LAI = Leaf area index (4.2), k = Constant governing the radiation extinction by the canopy (0.5)

#### **Diurnal Variation of Transpiration in HYDRUS-1D Model**

$$T_{p}(t) = 0.24\overline{T_{p}} \qquad t < 0.264d, t > 0.736d$$

$$T_{p}(t) = \overline{T_{p}} \sin\left(\frac{2\pi t}{1 \text{ day}} - \frac{\pi}{2}\right) \quad t \in (0.264d, 0.736d)$$

# **Validation of HYDRUS Model**



# **Validation of HYDRUS Model**

### **Modeled and Measured Transpiration**



Correlation Coefficient = 0.92 Nash–Sutcliffe efficiency (NSE)= 0.68

#### The model is able to reproduce sap flow values reasonably well

# Results: Applied Flux, Transpiration and Drainage Below Root Zone



### Drainage below the roots was 30.0 cm Applied top flux is exorbitantly high

# Results: Applied Flux, Transpiration and Drainage Below Root Zone



Drainage below the roots was 30.0 cm Applied top flux is exorbitantly high and need to be optimised

Sensitivity analysis is the study of the effect of the variation of input parameters on the output of the model

Partial derivative based analysisLocal or one-at-a-time (OAT) analysis

### **Global Sensitivity Analysis**

Covers all the input parameters including the affect generated due to the interaction of the parameters

More accurate and can be applied on complex and nonlinear models

### Variance based Sobol' method

- Variance based Sobol' method is a widely used algorithm for environmental models
- •The variance of the model output can be decomposed in terms of different fractions
- Each fraction represents the affect of a particular parameter and its interaction with other parameters
- The sensitivity of parameters are expressed in terms of sobol's sensitivity indices

### **Sobol Total Sensitivity Index**

$$S_{ti} = \frac{(1/2N)\sum_{j=1}^{N} \left[f(A)_{j} - f(A^{i}B)_{j}\right]^{2}}{(1/N)\sum_{j=1}^{N} \left[f(A)_{j}\right]^{2} - f_{o}^{2}}$$

$$S_{ti} = \frac{(1/2N)\sum_{j=1}^{N} \left[f(A)_{j} - f(A^{i}B)_{j}\right]^{2}}{(1/N)\sum_{j=1}^{N} \left[f(A)_{j}\right]^{2} - f_{o}^{2}}$$

$$f_o = (1/N) \sum_{j=1}^{N} f(A)_j$$
Hartman et al., 2017

A and B are the two set of random input parameter matrices

 $A^{i}B$  represents a matrix where all the columns are from matrix A except i<sup>th</sup> column (from matrix B)

N=5000, Nossent et al (2011)

The total number of simulations N(P+2) simulations = 35000 Saltelli (2002)

### **Steps of Sobol' Total Sensitivity Index Calculation**



Soil Hydraulic Parameters							
Parameter	Lower limit	<b>Upper limit</b>					
$\theta_r$	0.055	0.1					
$\theta_s$	0.38	0.5					
α	0.01	0.13					
n	1.25	2.3					
K <sub>s</sub>	6.0	355.0					

#### $P_i = P_{min} + (random \ value \times (P_{max} - P_{min}))$

a4 b4 c4 **d**44 e4

				4							В						
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	a3	b	3 c	3 đ	3	e3		a33	<b>b</b> 3	3	c33	<b>d</b> 33	3 e3	33			
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	ļ	18	5			A2B								Α	3B		
$\theta_r$	$\theta_s$	α	n	$K_s$		$\theta_r$	$\theta_s$	α	n	K	s	θ	r t	) <sub>s</sub>	α	n	$K_s$
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a4 b4 c4 d4 e44

### **Bootstrap Confidence Interval (BCI)**

#### Archer et al. (1997)

- •To estimate the accuracy of the total sensitivity indices (variation)
- •Randomly re-sampling (2500 samples with replacement) from the output space of each matrix (5000)
- •1000 values of total sensitivity indices calculated for each parameter



### **Total Sensitivity Index**



#### Drainage below root zone (-93 cm)

#### **Transpiration**

Parameter	Drainage below roo	t zone	Transpiration			
	Total Sensitivity (%)	BCI	Total Sensitivity (%)	BCI		
n	45	0.12	<10	0.043		
K <sub>s</sub>	36	0.085	58	0.27		
α	<1	< 0.001	60	0.24		
$\theta_s$	18	0.043	<1	< 0.001		
$\theta_r$	<1	< 0.001	<1	< 0.001		



- The cumulative transpiration ranged between 3.5-7.1 cm (50% variation)
- In clay type soil the oxygen stress is developed due to low hydraulic conductivity
- In case of sandy soil no water stress developed and transpiration happens at potential rate (T<sub>p</sub>) due to high water application rate
- Therefore the relative frequency distribution is left skewed and more than 60% simulations exhibit T<sub>p</sub>

### **Drainage Below the Root Zone**



- The cumulative drainage below root zone ranged between 26-54 cm
- High drainage observed towards sandy soil
- The relative frequency distribution is also left skewed

### **Irrigation Schedule Optimisation**



□ The water is getting stored in soil (23 cm)

□ Total 30 simulations -Decreasing applied Top flux by 0.1 cm/day

**Lesser decrease in transpiration** than the drainage below root zone

# **Irrigation Schedule Optimisation**

### **Optimization Criteria**

Changing Applied Irrigation
 Changing Irrigation interval
 Changing Initial soil moisture condition



Initial soil moisture 0.2



2 cm per day for 4 days then 25 days interval

Initial soil moisture 0.33



0.3 cm per day at every 12 days interval

# Conclusion

- **Good agreement** was achieved between HYDRUS-1D simulations and field measured sap flow
- The WUE for the present practice of flood irrigation was observed to be only 20%
- The GSA shows pore-size distribution index and saturated hydraulic conductivity has a major influence on the leakage below the root zone
- In contrast, the air-entry-pressure parameter and saturated hydraulic conductivity have a major influence on transpiration
- The initial conditions (Soil-water) play a significant role in calculating WUE
- Sensor based approach to trigger and control irrigation should be adopted for high WUE

# THANK YOU

Does anyone have any question?

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