

1 Article

2 Identification of phytoplankton blooms under the 3 index of Inherent Optical Properties (IOP index)

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19 **Abstract:** Phytoplankton blooms are sporadic events in time and isolated in space. This complex
20 phenomenon is produced by a variety of both natural and anthropogenic causes. Early detection of
21 this phenomenon, as well as the classification of a water body under conditions of bloom or non-
22 bloom, remains an unresolved problem. This research proposes the use of Inherent Optical
23 Properties (IOP) in optically complex waters to detect the bloom or non-bloom state of the
24 phytoplankton community. An IOP index is calculated from the absorption coefficients of the
25 colored dissolved organic matter (CDOM), the phytoplankton (φ) and the detritus (d), using the
26 wavelength (λ) 443 nm. The effectiveness of this index is tested in five bloom events in different
27 places and with different characteristics from Mexican seas: 1. Dzilam (Caribbean Sea, Atlantic
28 Ocean) a diatom bloom (*Rhizosolenia hebetata*), 2. Holbox (Caribbean Sea, Atlantic Ocean) a mixed
29 bloom of dinoflagellates (*Scrippsiella* sp.) and diatoms (*Chaetoceros* sp.), 3. Campeche Bay in the Gulf
30 of Mexico (Atlantic Ocean) a bloom of dinoflagellates (*Karenia brevis*), 4. Upper Gulf of California
31 (UGC) (Pacific Ocean) a diatoms bloom (*Planktoniella sol*) and 5. Todos Santos Bay, Ensenada (Pacific
32 Ocean) a dinoflagellates bloom (*Lingulodinium polyedrum*). The diversity of sites show that the IOP
33 index is a suitable method to determine the bloom conditions.

34 **Keywords:** Absorption coefficients, phytoplankton, detritus, CDOM, water quality, monitoring

35 1. Introduction

36 Phytoplankton blooms are sporadic events in time and isolated in space [1]. This complex
37 phenomenon is produced by a variety of both natural and anthropogenic causes [2]. The availability
38 of light and nutrients are key factors for its development [3]. These two factors concur during the
39 spring-summer period. At the beginning of this period the seasonal increase in daily irradiation
40 eliminates the light limitation, and the end of the thermal stratification supposes a supply of nutrients
41 thanks to the turbulent and convective mixing processes, which allows the phytoplankton to grow
42 rapidly [4]. However, phytoplankton blooms are not only limited to this period.

43 A bloom is the rapid growth of one or more species leading to an increase in the species' biomass [5].
 44 Different adjectives have been used to characterize the degree of negative impact of these blooms
 45 according to their characteristics and those of the causative species, such as toxic, noxious or harmful
 46 [6]

47 Identifying phytoplankton blooms has been the target of several research [7, 8, 9, 10]. Some research
 48 has focused on detecting changes in chlorophyll *a* fluorescence, changes in the composition of
 49 plankton species [9], or increases in nutrient levels [11]. Measuring blooms intensity has also been the
 50 subject of several research, such as continuous measurements of fluorescence and chlorophyll *a* [12]
 51 deviations in normal biomass variations [13], the ratio of two in situ optical measurements such as
 52 chlorophyll fluorescence (Chl *F*) and optical particulate backscattering (b_{bp}) [14], or satellite indices,
 53 such as the Maximum Chlorophyll Index (MCI) of the MERIS sensor [15]

54 Defining under which conditions an increase in phytoplankton biomass can be considered as a bloom
 55 is important to avoid an arbitrary use of the term bloom [7, 16, 4]. This research proposes the use of
 56 Inherent Optical Properties (IOP), specifically the absorption coefficient, as an indicator that a
 57 phytoplankton community has passed into a bloom condition.

58 The absorption coefficient $a(\lambda)$ characterizes light absorption properties in the aquatic environment.
 59 Light absorption in natural waters is attributable essentially to four components: water, colored
 60 dissolved organic matter, photosynthetic biota and inorganic particles [17]. Thus, $a(\lambda)$ can be
 61 expressed as:

$$62 \quad a(\lambda) = a_w(\lambda) + a_{cdom}(\lambda) + a_p(\lambda) \quad (1)$$

63 Where the subscripts *w*, *cdom* and *p* represent water, colored dissolved organic matter (CDOM) and
 64 particulate matter, respectively. This particulate material consists of phytoplankton (φ) and detritus
 65 (non-algal particles) (*d*), thus, $a_p(\lambda) = a_\varphi(\lambda) + a_d(\lambda)$ [18].

66 Seawater components present a typical spectrum of light absorption, which means that they absorb
 67 light with a preference for certain wavelengths in the visible (400 to 700 nm) or ultraviolet (250 to 400
 68 nm) [17]. Optically pure water $a_w(\lambda)$ absorbs light with a preference for red in the electromagnetic
 69 spectrum of 750 to 800 nm. Phytoplankton has an absorption spectrum $a_\varphi(\lambda)$ characterized by two
 70 peaks located in the 440 and 675 nm spectrum, which are related to chlorophyll *a* absorption. Detritus
 71 $a_d(\lambda)$ and CDOM $a_{cdom}(\lambda)$ absorb with an exponential increase towards shorter wavelengths, with
 72 the most significant absorption towards the UV spectrum between 250 and 400 nm [19]. In optically
 73 complex waters, such as coastal and inland waters, the optical properties are determined by the
 74 combination of these water components in varying proportions [20].

75 [19], developed the IOP index with the objective of identifying phytoplankton blooms. This index is
 76 calculated from the absorption coefficients of the colored dissolved organic matter (CDOM), the
 77 phytoplankton (φ) and the detritus (*d*), using the wavelength (λ) 443 nm, and the relationship with
 78 chlorophyll *a* concentration and phytoplankton abundance is analyzed.

79 This research proposes the use of Inherent Optical Properties (IOP) in optically complex waters to
 80 detect the bloom or non-bloom state of the phytoplankton community, as well as it is an active or a
 81 decaying bloom. The objective is to test the effectiveness of IOP index in bloom events in different
 82 coastal areas with distinctive characteristics.

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86 2. Materials and Methods

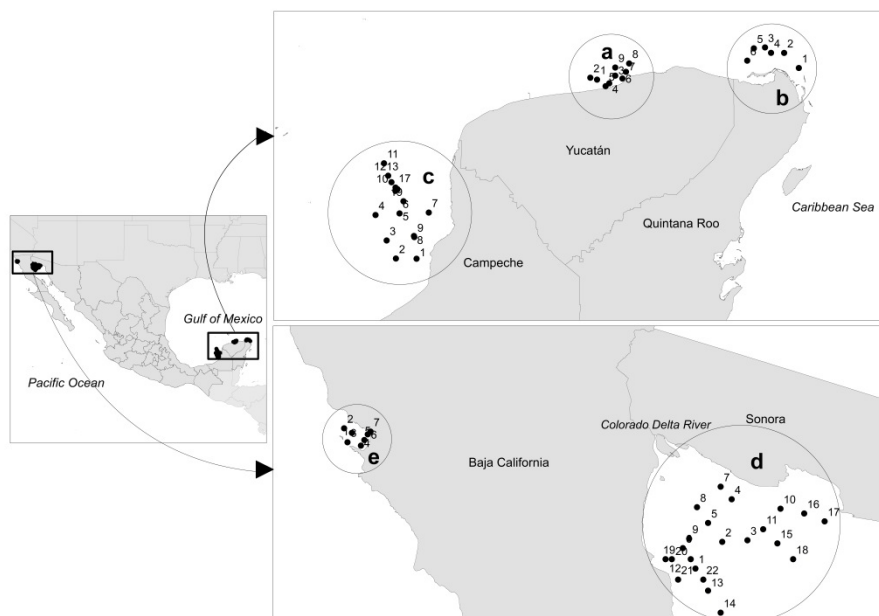
87 2.1. Study area

88 The study area are well-known coastal areas of Mexico with distinctive characteristics where bloom
89 events have been observed recurrently (Fig. 1). These areas are:

90 Area 1, three coastal areas in the Yucatán Peninsula: Dzilam de Bravo (Dzilam for short) in the
91 Yucatan state (Fig. 1a), Holbox in the Quintana Roo state (Fig. 1b), and Campeche Bay in the
92 Campeche state (Fig. 1c). This Peninsula is a karstic region, characterized by minimal soil cover and
93 rapid infiltration of rain water, with the consequent high vulnerability of aquifer pollution [21, 22].
94 The rainy season occurs from June through December with minimal rainfall occurring during the rest
95 of the year. The unconfined Yucatán aquifer has submarine groundwater discharges (SGD) that can
96 threat coastal ecosystems [23, 22]. SGD has been linked to eutrophication and harmful algal blooms
97 [23]. According to [24], the Yucatán coastal aquifer is a triple porosity system, where the flow of
98 groundwater takes place mainly through interconnected cave systems and fractures, and drains
99 inland catchment mainly through coastal springs. In recent years, intense coastal development is
100 taking place within the Caribbean, due to tourism, which increases risk of aquifer pollution. This
101 development is particularly fast in the eastern coast of the Yucatan Peninsula (Quintana Roo state).
102 Both Yucatán and Quintana Roo state coastal waters are influenced by waters of the Caribbean Sea
103 and the Gulf of Mexico [25]. Campeche state coastal water is influenced by the current system of
104 Yucatan/Lazo/Florida [26]. This region has a predominantly cyclonic circulation [27], caused by the
105 wind effort [28], and by an upwelling on the north coast of the Yucatan Peninsula [29].

106 Area 2, the Upper Gulf of California (UGC). The Gulf of California is a semi-enclosed sea in the
107 Eastern Pacific. The UGC is located in the Northern Gulf of California, where the Sonora and Baja
108 California states coasts intersect at a 60° angle [30]. It is considered as one of the most biologically
109 productive marine regions [31, 32], with peak chlorophyll *a* concentrations of 18.2 mg m⁻³ and
110 averages of 1.8 mg m⁻³ between 1997 and 2007 in coastal waters near the delta [33]. This high
111 productivity is due to a complex mix of factors, including: coastal upwelling, wind-driven mixing,
112 extreme tidal mixing and turbulence, thermohaline circulation, coastal-trapped waves, regular
113 sediment resuspension, and, to a lesser extent, agricultural runoff, released nutrients from erosion of
114 ancient Colorado River Delta sediments and groundwater discharges [31, 34]. After the construction
115 of the Hoover and Glen Canyon dams in the USA in 1935 and 1964, the Colorado river only discharges
116 variable and insignificant surface water-flows occasionally into the Gulf of California [34].

117 Area 3, Todos Santos Bay (TSB), is a semi-enclosed bay, adjacent to the Pacific Ocean, within the
118 upwelling zone of the Baja California peninsula (Mexico). This area is influenced by the California
119 Current System (CCS), which produces coastal upwelling along the coast of the Baja California
120 peninsula. This is a phenomenon with a marked seasonality caused by the prevailing winds from the
121 northwest, which tend to be more intense during the spring and summer months [35, 36, 37]. Two
122 water masses integrate the CCS, the California Current (CC), a year-round equatorward surface flow,
123 which transports Subarctic Water (SAW), characterized by low salinity, and the California
124 Undercurrent (CU), a poleward subsurface (100–400 m) flow that transports Equatorial Subsurface
125 Water (ESsW), characterized by relatively high salinity, high nutrient concentration, and low
126 dissolved oxygen content, according to [38] description. SAW are mainly important during winter
127 and spring, while ESsW appear at the end of summer and autumn [39]. In addition to the described
128 seasonal variability, the El Niño-Southern Oscillation (ENSO) induces oceanographic changes in the
129 region off Baja California at an interannual scale [39]. Altogether, these factors control primary
130 productivity which is characteristically high [35, 40]. Dinoflagellate algal blooms (DABs) events in
131 this area have increased considerably in extension and frequency over the past two decades [41].



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133 **Figure 1.** Sampling stations. a) Dzilam de Bravo (Yucatan), b) Holbox (Quintana Roo), c) Campeche
 134 Bay (Campeche), d) Upper Gulf of California (Baja California and Sonora) and e) Todos Santos Bay
 135 (Baja California).

136 2.2. Collection of samples

137 Water samples were taken in Mexico coastal waters at the stations shown in Fig.1. Samples were
 138 taken in four field campaigns, two on year 2011 and two on 2017, during reported bloom events.

139 Dzilam (Yucatán) and Holbox (Quintana Roo) samples were collected between the 27th and the 30th
 140 of August 2011 (9 and 6 samples respectively). All the Dzilam and Holbox stations were sampled at
 141 surface (1.5 m), stations were selected based on reports of fishermen on fish mortality and patches of
 142 discolored water. Campeche Bay (Campeche) samples were collected between the 22nd and the 24th
 143 of September 2011 (19 samples). Campeche Bay was also sampled at surface (1.5 m), except for
 144 stations number 13 and 16 which were sampled at 15 m. The campaign was conducted in response
 145 to a phytoplankton bloom reported by various local, state and federal public health institutions in
 146 Campeche. The Todos Santos Bay (TSB) in Ensenada (Baja California) was sampled on June 2, 2017
 147 (7 samples) during the second week of a bloom event that lasted three weeks. This event was
 148 characterized by the bioluminescence observed during all the nights that lasted. TSB was also
 149 sampled at surface (0.5 m). Stations 5, 6 and 7 were taken on the reddish patch that distinguished
 150 itself from the rest of the bay water.

151 These data were collected in small vessels where the samples were taken manually and stored in
 152 Nalgene dark bottles of high density polyethylene (HDPE) until processing in the laboratory. For the
 153 CDOM samples were collected in amber glass bottles and refrigerated until laboratory processing.

154 Sampling of the Upper Gulf of California (UGC), was carried out from February 23 to March 3, 2017,
 155 on the research vessel "Tecolutla" of the Mexican Navy during the oceanographic cruise "Vaquita
 156 Marina 2017" (22 samples). Samples were taken with Niskin bottles attached to a rosette, and
 157 immediately processed in the vessel's laboratory. Sampling depth was at the chlorophyll maximum
 158 fluorescence (10 to 40 m). The chlorophyll maximum was measured with an ECO FLNTU fluorimeter
 159 coupled to a CTD SB 19 Plus. During the oceanographic cruise color patches were detected in the
 160 water, on this basis it was decided to take samples.

161 In each study area the samples were collected inside and outside the patches with evidence of a
 162 bloom, in order to be able to capture the variability that exists in a parcel of water, and to better define
 163 the baseline or mean of each campaign.

164 2.3 Absorption coefficients determination

165 The CDOM samples were filtered using a 0.2 µm pore membrane filter (Nuclepore™) and processed
 166 according to the methodology of [42]. The CDOM absorption coefficient, $a_{cdom}(\lambda)$, was measured in
 167 the wavelength range of 250 to 800 nm in a 10 cm long quartz cuvette using Milli-Q water as reference.

168 Particulate matter absorption coefficient was determined using the methodology of [42]. A volume
 169 of seawater of 0.5 to 2 L, depending on the particle load, was filtered from water stored in Nalgene
 170 bottles, with Whatman GF/F glass fiber filters 25 mm in diameter and 0.7 µm in size of pore. The
 171 particulate matter absorption coefficient, ($a_p(\lambda)$), was measured in the wavelength range of 400 to
 172 800 nm. Then, the filters are immersed in methanol to depigment the filter and obtain the detritus
 173 coefficient absorption, $a_d(\lambda)$. The phytoplankton absorption coefficient, $a_\phi(\lambda)$, was calculated by
 174 subtracting $a_d(\lambda)$ from $a_p(\lambda)$.

175 The 2011 samples were read with a Perkin-Elmer Lambda 18 spectrophotometer, and the 2017
 176 samples were read with a Cary 100 UV-Visible spectrophotometer.

177 A non-parametric one-way analysis of variance (Kruskal–Wallis) was performed to statistically assess
 178 variations in the absorption coefficients. The water absorption coefficient of phytoplankton, detritus
 179 and CDOM for each sampling area was compared.

180 2.4 IOP index determination

181 The IOP index was determined according to [19] following the next steps. Firstly, the absorption
 182 coefficients ($a_{cdom}(443)$, $a_d(443)$, $a_\phi(443)$) were standardized, and a principal component
 183 analysis (PCA) was performed to explore associations between the sampled stations. Then, samples
 184 were classified as bloom or non-bloom using a factorial analysis [43]. Finally, the IOP index was
 185 calculated based on the first standardized empirical orthogonal function (SEOF₁) [19] according to
 186 equation (2).

$$187 \quad IOP_{index} = -1[(b_{1,1} * Z a_{phy,443}) + (b_{1,2} * Z a_{CDOM,443}) + (b_{1,3} * Z a_{d,443})] \quad (2)$$

188 The coefficients $b_{1,1}$, $b_{1,2}$ y $b_{1,3}$ are the eigenvalues resulting from the PCA, while $a_{phy,443}$,
 189 $a_{CDOM,443}$ and $a_{d,443}$ are the values obtained from the Pearson correlation matrix between the
 190 absorption coefficients. To describe the stages of a phytoplankton bloom, [19], interpreted the values
 191 of the IOPs index as: 1) values in the interval (-1,1) show an average value and represent non-bloom
 192 conditions; 2) values in the interval (1, 2) are above the average and represent decaying bloom
 193 conditions, and 3) values higher than 2 are anomalous and indicate active bloom conditions.

194 2.5 Phytoplankton characterization

195 The blue/red ratio (B/R) is an index that allows to characterize the dominant phytoplankton size [44,
 196 45, 46, 47, 19]. It is calculated as expressed in equation (3):

$$197 \quad B/R = \frac{a_{phy,443} (440)}{a_{phy,443} (675)} \quad (3)$$

198 If the B/R is >3.0 , dominance of picophytoplankton ($<2 \mu\text{m}$) is implied. If the ratio is <2.5 , dominance
 199 of microphytoplankton ($>20 \mu\text{m}$) is implied. Ratios between 2.5 and 3.0 indicate that there is no
 200 dominance of a particular group and is identified as mixed bloom

201 Some representative samples of each sampling were analyzed by microscopy to identify the main
 202 blooming specie and/or genus. Samples were preserved in 125 ml bottles in a neutral lugol solution
 203 with a sodium acetate base in a 1:100 ratio. The samples were stored in dark and cold conditions until
 204 their identification. The Dzilam, Holbox, and Campeche samples were identified at the Florida Fish
 205 and Wildlife Conservation Commission (FWC). Phytoplankton identification was performed using
 206 an inverted Olympus IX71 microscope following a modified method of Utermöhl [48]. In the case of
 207 the UGC and TSB samples, the same method was performed using phase contrast microscopy with
 208 a microscope Bausch and Lomb. [49, 50, 51, 52], were used as taxonomic references.

209 For Dzilam, Holbox and Campeche the chlorophyll *a* concentration was determined fluorometrically
 210 on methanol extracts following the method of [53], using a Turner Designs 10-AU field fluorimeter.

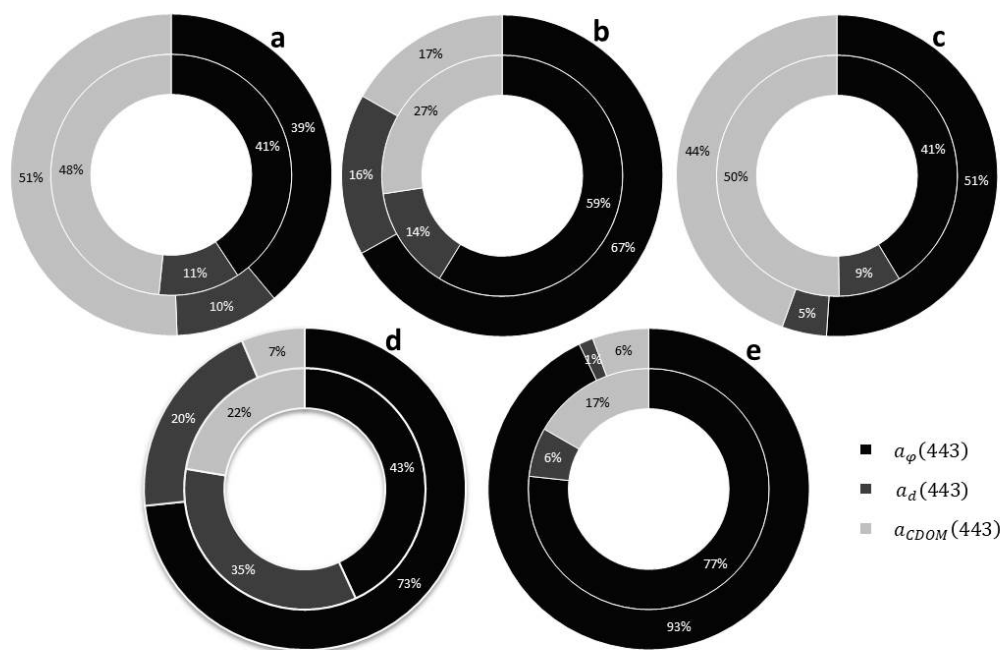
211 3 Results and discussion

212 In Table 1, we summarized the main characteristics of each bloom event. For each sampling
 213 campaign, we studied the contribution of each water component absorption coefficient (colored
 214 dissolved organic matter ($a_{cdom}(443)$), phytoplankton ($a_p(443)$) and detritus ($a_d(443)$) to the
 215 global absorption coefficient $a(443)$ at 443 nm. In Figure 2, the inner circumference shows the
 216 average contribution of each absorption coefficient to $a(443)$ for each sampling campaign.

217 **Table 1.** Characterization of bloom events. The absorption coefficient that contributes most to the
 218 Inherent Optical Properties (IOP) is underlined for each sampling campaign.

Blooms	# Samples	Samples in active bloom	Dominant phytoplakton species	Dominant population size	Proportion of IOP at the station with Bloom (%)		
					$a_p(443)$	$a_d(443)$	$a_{cdom}(443)$
1 Dzilam	9	1	<i>Rhizosolenia hebetata</i> <i>Scrippsiella</i> sp	Microphytoplankton	39	10	<u>51</u>
2 Holbox	6	1	<i>Chaetoceros</i> sp <i>Rhizosolenia hebetata</i>	Mixed community	<u>67</u>	16	17
3 Campeche	19	4	<i>Karenia brevis</i>	Microphytoplankton	<u>51</u>	5	44
4 UGC	23	1	<i>Planktoniella sol</i>	Picophytoplankton	<u>73</u>	20	7
5 TSB	7	1	<i>Lingulodinium polyedrum</i>	Microphytoplankton	<u>93</u>	1	6

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Figure 2. Contribution of each absorption coefficient ($a_{\phi}(443)$, $a_d(443)$ and $a_{cdom}(443)$) to $a(443)$ for each sampling area. The inner circumference shows the average contribution of each absorption coefficient to $a(443)$ for each sampling campaign. the outer circumference represents the average value of sampling points classified as active bloom according to the IOP index. **a** Dzilam de Bravo **b** Holbox. **c** Campeche Bay. **d** Upper Gulf of California. **e** Todos Santos Bay.

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In Dzilam, colored dissolved organic matter (CDOM) was the major contributor to $a(443)$. $a_{cdom}(443)$ represented the 48% of total absorption, followed by $a_{\phi}(443)$ with 41% and $a_d(443)$ with 11% (Fig. 2a). In Holbox, phytoplankton was the major contributor to $a(443)$. $a_{\phi}(443)$ represented the 59% of water absorption, followed by $a_{cdom}(443)$ with 27% and $a_d(443)$ with 14% (Fig. 2b). In Campeche Bay, as in Dzilam, the dominant absorption component was CDOM, $a_{cdom}(443)$ was 50%, followed very close by phytoplankton $a_{\phi}(443)$ was 41% of $a(443)$, and a minor contribution of detritus ($a_d(443)$ was 9%) (Fig. 2c). In the Upper Gulf of California the highest contribution was from phytoplankton ($a_{\phi}(443)$ was 43%), followed by detritus ($a_d(443)$ was 35% of $a(443)$), and CDOM ($a_{cdom}(443)$ was 22%) (Fig. 2d). In Todos Santos Bay (TSB), as in Holbox, phytoplankton represented the highest absorption percentage ($a_{\phi}(443)$ was 77%). However, in TSB the contribution of CDOM and detritus is characteristically low (17% and 6% respectively).

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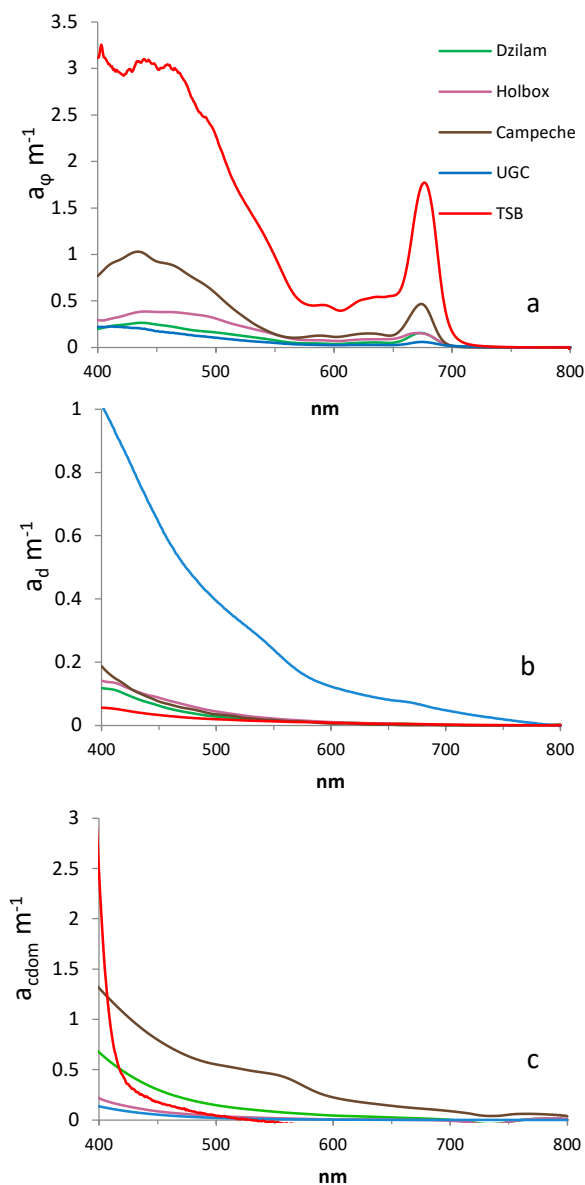
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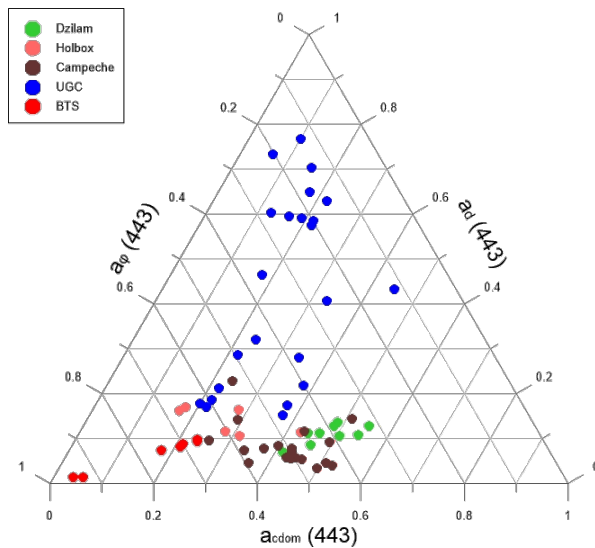
In Fig. 3, the phytoplankton, detritus and CDOM absorption spectrum of all sampling campaigns are compared. The phytoplankton absorption coefficient, $a_{\phi}(\lambda)$, was significantly higher in TSB than in other sampling areas ($P < 0.05$ for $a_{\phi}(443)$). No significant differences were observed between Dzilam and Campeche Bay ($P > 0.05$ for $a_{\phi}(443)$). The lowest $a_{\phi}(\lambda)$ values were observed in the UGC. The detritus absorption coefficient, $a_d(\lambda)$, was significantly higher in the UGC than in all the other studied areas ($P < 0.05$ for $a_d(443)$). No significant differences were observed between the Yucatan Peninsula areas (Dzilam, Holbox and Campeche), nor with TSB ($P > 0.05$ for $a_d(443)$). The CDOM absorption coefficient, $a_{cdom}(\lambda)$, was significantly higher in Dzilam and Campeche Bay than in other areas ($P < 0.05$ for $a_d(443)$).



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247 **Fig. 3** Absorption coefficients ($a(\lambda)$): a) phytoplankton, b) detritus and c) colored dissolved organic matter (CDOM) of sampling points in active bloom for each sampling campaign (Dzilam, Holbox,
 248 Campeche Bay, Upper Gulf of California (UGC) and Todos Santos Bay (TSB))
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250 In Figure 4, it is represented the spectrum of absorption of each seawater component
 251 (phytoplankton, detritus and colored dissolved organic matter) for all the sampling points
 252 (Dzilam, Holbox, Campeche Bay, Upper Gulf of California and Todos Santos Bay). This
 253 graphical representation allowed us to compare the different study areas. In general terms,
 254 the most important components were phytoplankton ($a_{\phi}(443)$) and CDOM
 255 ($a_{cdom}(443)$). In this graph, we observed the significantly higher importance of detritus in
 256 the UGC. This detritus contribution is much more important near de Colorado River and
 257 decreases southward.



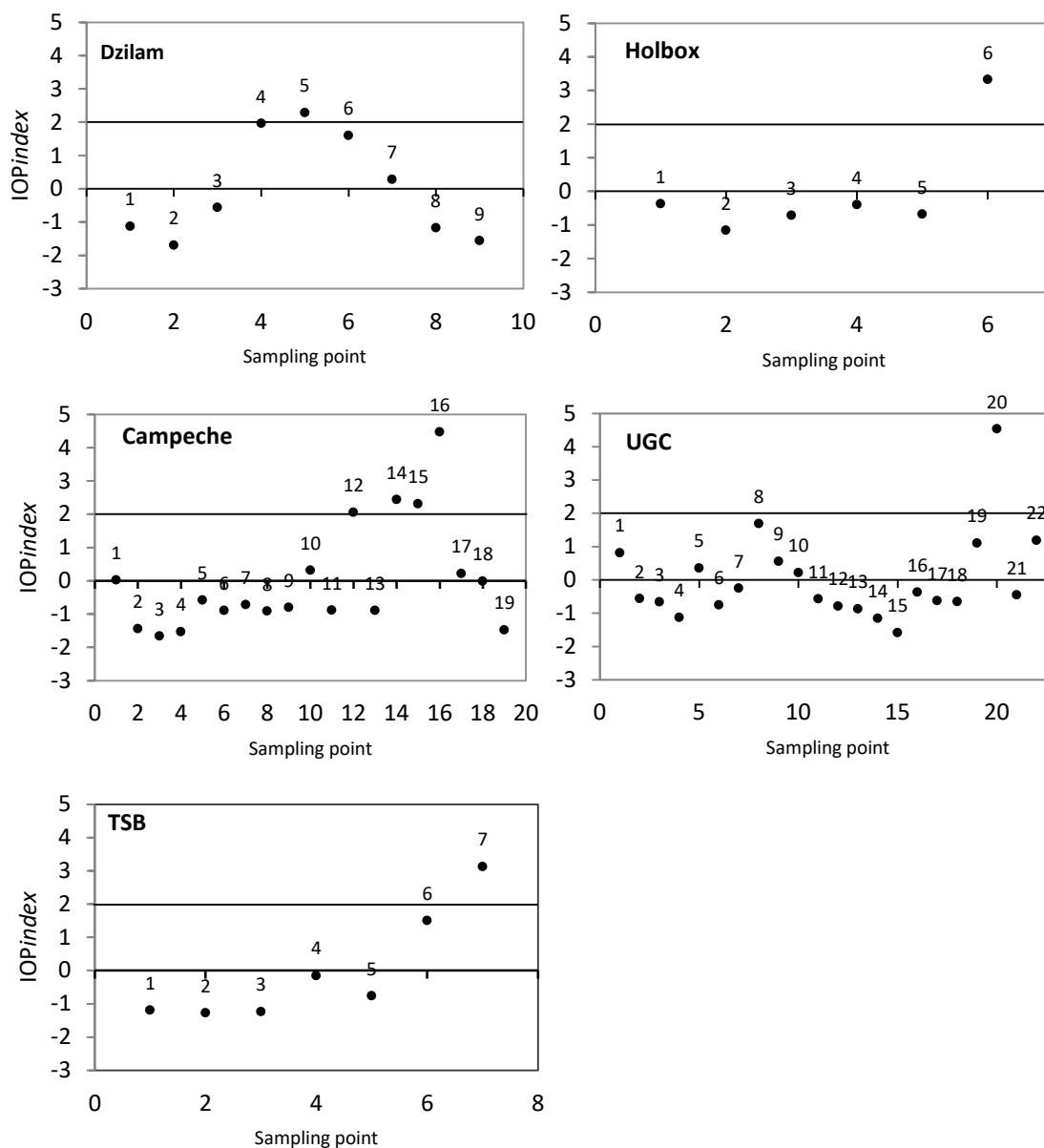
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Fig. 4 Triangular diagram used to classify sampling points according to the contribution to $a(443)$ of each component: phytoplankton ($a_p(443)$), colored dissolved organic matter ($a_{cdom}(443)$) and detritus ($a_d(443)$).

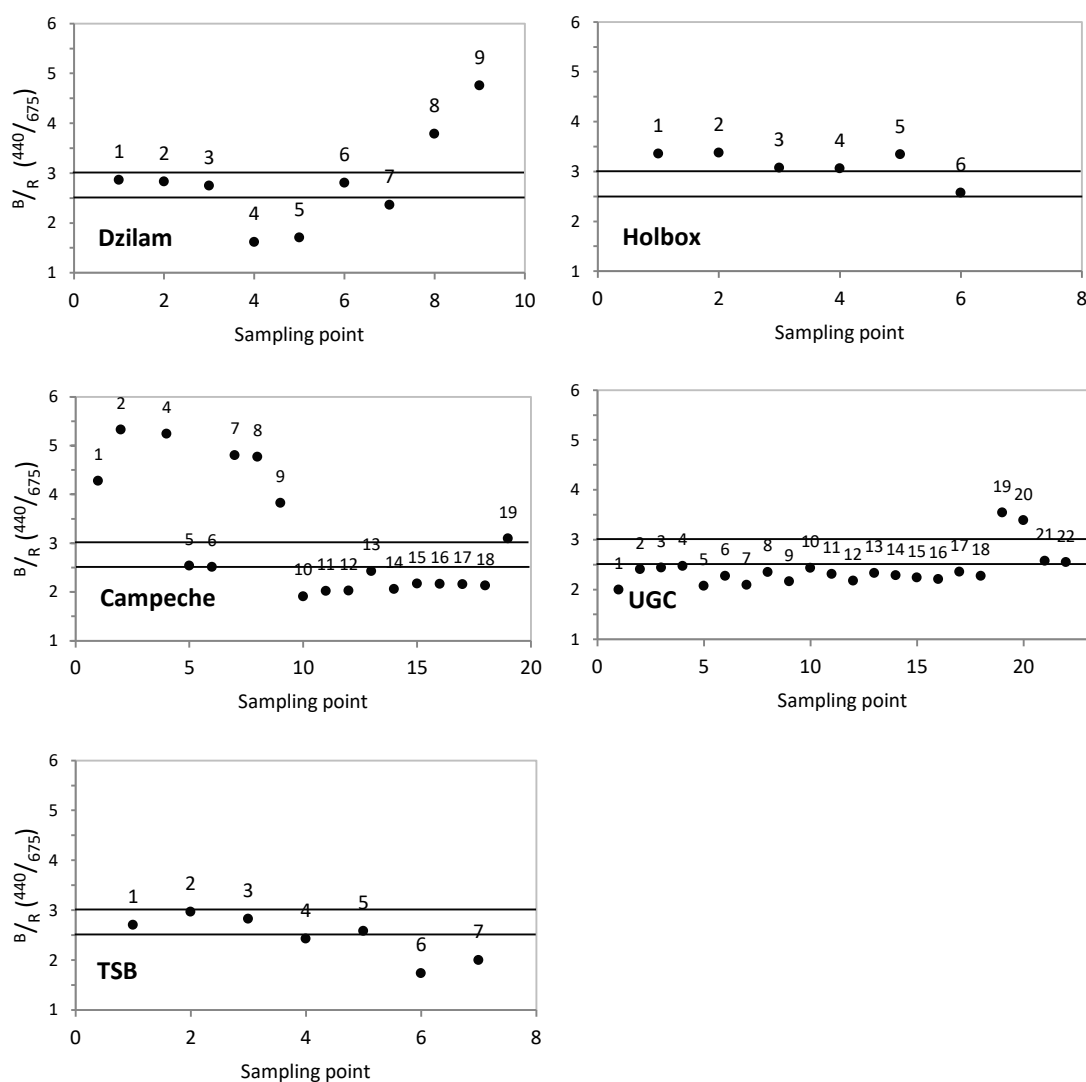


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263 **Fig. 5.** IOP index results for each sampling campaign and sampling point. From top to bottom and
 264 from left to right: Dzilam, Holbox, Campeche Bay, Upper Gulf of California (UGC) and Todos Santos
 265 Bay (TSB)

266 The IOP index was calculated from the absorption coefficients for each sampling area and sampling
 267 point. IOP index results are represented graphically in Fig. 5. In Fig. 2 the outer circumference
 268 represents the average value of sampling points classified as active bloom according to the IOP index.
 269 In Dzilam, sampling points 4 and 6 had a value in the interval (1, 2), meaning that they were above
 270 the sampling area average and in decaying bloom conditions. However, only sampling point 5 was
 271 above two and in active bloom conditions. In Fig. 2a, we observed that the contribution of each
 272 absorption coefficient to $a(443)$ in sampling point 5 is similar to the sampling campaign average. In
 273 Holbox, only sampling point 6 was above an IOP index value of two (Fig. 5), and thus in active bloom
 274 conditions. In Fig. 2b, we observed a higher contribution of phytoplankton to $a(443)$ than the
 275 average value of the sampling campaign ($a_p(443)$ of 67% in sampling point 6 compared with 59%
 276 average value). The lower average contribution of phytoplankton when considering all sampling
 277 points was related with a higher CDOM contribution in non-bloom conditions. In Campeche Bay,
 278 sampling points 12, 14, 15 and 16 were in active bloom conditions (Fig. 5). Sampling point 16 showed

279 the highest anomaly; this sample was collected at 15 m depth. In Fig. 2c, as in Holbox, we observed a
 280 higher contribution of phytoplankton to $a(443)$ than the average value of the sampling campaign
 281 ($a_{\phi}(443)$) of 51% in sampling point 16 compared with 41% average value). The lower average
 282 contribution of phytoplankton was also related with a higher CDOM contribution in non-bloom
 283 conditions. In the Upper Gulf of California (UGC), sampling points 8, 19 and 22 were in decaying
 284 bloom conditions (IOP index value higher than one and lower than two), while sampling station 20
 285 was in active bloom conditions according to IOP index (Fig. 54). In Fig. 2d, we observed that, as in
 286 Holbox and Campeche Bay, the contribution of phytoplankton to $a(443)$ was higher than the
 287 average ($a_{\phi}(443)$) of 73% in sampling point 20 compared with 43% average value). In Todos Santos
 288 Bay, sampling point 6 was under decaying bloom conditions, while sampling point 7 was in active
 289 bloom conditions. As in Holbox and Campeche Bay, we noticed a higher contribution of
 290 phytoplankton to $a(443)$ than the average value of the sampling campaign ($a_{\phi}(443)$) of 93% in
 291 sampling point 7 compared with 77% average value). $a_{cdom}(443)$ and $a_d(443)$ contribution was
 292 even lower than average.



293
 294 **Fig. 6.** B/R index for each sampling campaign and sampling point. From top to bottom and from left
 295 to right: Dzilam, Holbox, Campeche Bay, Upper Gulf of California (UGC) and Todos Santos Bay
 296 (TSB).

297 In order to characterize the phytoplankton community, the blue/red ratio (B/R) is graphically
 298 represented in Fig. 6. B/R values higher than 3 reveal a community dominated by picophytoplankton;

299 B/R values lower than 2.5 reveal microphytoplankton (>20 μ m) dominance; and B/R values between
300 2.5 and 3.0 indicated mixed community. In Dzilam, microphytoplankton dominated in the active
301 bloom sampling point 5 ($B/R = 1.71$) (Fig. 6). According to the microscope taxonomic analysis, the
302 dominant specie was the diatom *Rhizosolenia hebetata*. In Holbox, the B/R ratio in active bloom point
303 6 was 2.57 (Fig. 6), thus a mixed picophytoplankton and microphytoplankton community was
304 observed. This was corroborated by microscope taxonomic analysis that identified the dinoflagellate
305 *Scrippsiella* sp., and the diatoms *Chaetoceros* sp. and *Rhizosolenia hebetata*. In Campeche Bay, B/R was
306 lower than 2.5 in all active bloom conditions points (Fig. 6), so microphytoplankton was dominant.
307 The dinoflagellate *Karenia brevis* was identified by microscopy as the dominant specie. In the UGC,
308 B/R was below 2.5 in nearly all the sampling stations (Fig. 6). However, in sampling point 15, B/R
309 was 2.59 pointing out a mixed community. The diatom *Planktoniella sol* was identified by microscopy.
310 In Todos Santos Bay, B/R was below 2.5 in sampling point 7 (active bloom conditions) (Fig. 6), thus
311 indicating microphytoplankton dominance. The most abundant specie in this point was the
312 dinoflagellate *Lingulodinium polyedrum*.

313 Dzilam (Yucatan), Holbox (Quintana Roo), and Campeche Bay (Campeche) (Fig. 1a, b & c) are located
314 in the karstic Yucatan Peninsula [54]. This region is characterized by rapid rain water infiltration into
315 the groundwater system, and nearly no surface runoff [21, 25]. Due to its hydrological characteristics
316 the lowest absorption coefficient is the detritus one ($a_d(443)$) is 11%, 14% and 9% respectively in each
317 area) (Fig. 3), as there is no relevant detritus source, no river runoff (the nearest one are located in
318 south Campeche, far from the sampling area located in north Campeche). The climate of the region
319 is characterized by three seasons associated with rainfall patterns: the dry season (March to May), the
320 rainy season (June to October) and the northern wind season [55]. In this region, submarine
321 groundwater discharges (SGD) play a significant role in driving the nutrient stoichiometry (N:Si:P
322 ratio) in receiving waters, which is a key factor in phytoplankton assemblages. SGD are an important
323 source of nitrogen, particularly NO_3^- , during the wet season (June to October), the high N:P ratio in
324 SGD can drive phosphorus limitation in the nearshore environment [23]. SGD are also rich in silica,
325 which can conduct to diatom growth. Several studies have corroborated low salinity groundwater as
326 an important source of nutrients in the Yucatan, specifically NO_3^- and silica, and have linked SGD to
327 harmful algal blooms [23]. According to [55] the HAB events in the state of Yucatan have been
328 reported almost every year since 2001, covering an approximate area of 6000 km².

329 Our sampling was developed during the August–December 2011 large scale pelagic HAB event. This
330 event started in Dzilam and tended to move westward along the northern Yucatan coast [54]. In
331 Dzilam, the dominance of the diatom *Rhizosolenia hebetata* can be explained by the input of silica from
332 near near springs (cenotes). [56] observed the the maximum chlorophyll *a* concentrations on August
333 8 and 30. Our sampling was performed on August 27. So, the degradation of phytoplankton cells
334 from the previous peak may explain the high contribution of CDOM absorption coefficient (48% on
335 average). The sampling point identified as in active bloom conditions according to the IOP index had
336 significantly higher chlorophyll *a* levels, 12.5 mg m⁻³, that points in non-bloom conditions, 3.1 mg m⁻³
337 on average.

338 In Holbox, diatoms were also dominant, *Chaetoceros* sp. and *Rhizosolenia hebetata*, but dinoflagellates
339 of *Scrippsiella* sp. were also abundant. Both *Chaetoceros* sp. and *Scrippsiella* sp. were also observed in
340 Dzilam during this HAB event according to [56]. The characteristics springs (cenotes) of the Quintana
341 Roo state could have supplied the silica needed for this sustained diatom bloom. Also in this

342 sampling campaign, the sampling point identified as in active bloom conditions according to the IOP
343 index had significantly higher chlorophyll *a* levels, 12.5 mg m⁻³, that points in non-bloom conditions,
344 2.2 mg m⁻³ on average.

345 In Campeche Bay, the blooming specie was identified as the dinoflagellate *Karenia brevis*. Again, in
346 this sampling campaign, the point in active bloom conditions according to the IOP index had
347 significantly higher chlorophyll *a* levels, 33.2 mg m⁻³, that points in non-bloom conditions, 7.0 mg m⁻³
348 on average. The CDOM absorption coefficient, $a_{cdom}(443)$, was as high as in Dzilam (higher than
349 in all other our study areas) (Fig. 3). Our sampling was performed on September 22, 2011. So, the
350 high CDOM values could be explained by the degradation of accumulated phytoplankton cells
351 during August and September. This region is influenced by the current system of
352 Yucatan/Lazo/Florida [26]. It is important to say that even under very high CDOM values, the IOP
353 index was able to distinguish an active phytoplankton bloom.

354 The Upper Gulf of California (UGC) and Colorado River Delta (CRD) area, is a region of sediment
355 re-suspension characterized by high detritus levels, low light extinction coefficient values (-0.05 m^{-1})
356 and high sedimentary loads (maximum values of 8 g/L) [30]. So, we expected the highest detritus
357 absorption coefficient ($a_d(\lambda)$) observed. It is remarkable that, also under very high detritus levels, the
358 IOP index was able to distinguish an active phytoplankton bloom.

359 In Todos Santos Bay (TSB), the most abundant species during our study was the dinoflagellate
360 *Lingulodinium polyedrum*. [41] have reported an increase in dinoflagellate algal blooms (DABs), with
361 *Lingulodinium polyedrum* as the dominant species, over the past few years in coastal areas off Baja
362 California. Our sampling was developed on June 2, 2017, that is late spring, when *L. polyedrum* blooms
363 usually occur in this area [41]. This blooms have been related with a increases in irradiance, daylight
364 hours, temperatures between 17 and 23°C, stratification of the water column and formation of a
365 seasonal surface thermocline [57]. These blooms are favoured by the convergence of surface currents
366 and winds, which induce the transport of cells that tend to concentrate near the surface and toward
367 the coast [41, 58]. This bloom presented the highest phytoplankton absorption coefficient ($a_p(\lambda)$)
368 observed in our study (Fig. 3).

369 4. Conclusions

370 The selected study areas have allowed us to apply the IOP index within the wide variability of coastal
371 waters, optically complex waters. Within this variability, we found areas with dominance of detritus
372 or CDOM, despite the samplings were developed in areas with observed phytoplankton blooms. The
373 IOP index was able to discern sampling points in active bloom conditions from points in decaying
374 bloom conditions. In the Yucatan region, the IOP index distinguished points in active bloom from
375 points with high CDOM due to phytoplankton cell degradation from previous bloom. Also, the IOP
376 index has been proved useful to distinguish phytoplankton blooms from the natural variability of
377 one area. In the case of the UGC, typical high detritus levels produce high absorption coefficient,
378 which is not related with phytoplankton blooms. The IOP index was able to identify points in active
379 bloom conditions from points with high detritus load.

380 To be able to distinguish a phytoplankton bloom from natural variability it is important regular
381 monitoring. The inherent optical properties play a key role for correctly identifying phytoplankton

382 blooms, but are highly variable in complex coastal waters. Different coastal areas have different
383 baseline values that should be defined to be able to detect anomalous events. Thus, the measurement
384 of absorption coefficients should be considered in coastal waters monitoring programs. The use of
385 remote sensing can help to define IOPs from satellite reflectances, $R_{rs}(\lambda)$, and to build a baseline at
386 a lower cost. Further research is needed to test if contrasting in situ IOPs measures, to a baseline
387 calculated by remote sensing, through the IOP index, is also able to correctly identify active
388 phytoplankton blooms.

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