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# Post-fire effect modeling for the permafrost zone in Central Siberia on the basis of remote sensing data

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13 Abstract: The increasing trend of larch forests burning in the permafrost zone (60–65° N, 95–105°E) 14 is observed in Siberia. Up to 10-15% of entire larch forests were damaged by wildfire during the 15 last two decades. Current research analysed the reflectance and thermal anomalies of the 16 post-pyrogenic sites under the conditions of permafrost. Studies are based on long-term Terra, 17 Aqua/MODIS (Moderate Resolution Imaging Spectroradiometer) survey for 2006–2018. We used IR 18 thermal range data of 10.780-11.280 microns (MOD11A1 product) and we evaluated NDVI from 19 MOD09GQ product as well. The averaged temperature and NDVI dynamics were investigated in 20 total for 50 post-fire plots under different stages of succession (1, 2, 5 and 10 years after burning) in 21 comparison with non-disturbed vegetation cover sites under the same conditions. We recorded 22 higher temperatures (20-47% higher than average background value) and lower NDVI values 23 (9–63% lower than non-disturbed vegetation cover) persisting for the first 10 years after the fire. 24 Under conditions of natural restoration thermal anomalies of the ground cover remained 25 significant for more than 15 years, which was reflected on long-term satellite data and confirmed 26 by ground-based measurements. To estimate impact of thermal anomalies on soil temperature and 27 thawed layer depth we used the Stefan's solution for the thermal conductivity equation. According 28 to results of numerical simulation, depth of the seasonal thawed layer could increase more than 29 20% in comparison with the average statistical norm under the conditions of excessive heating of 30 the underlying layers. This is a significant factor in the stability of Siberian permafrost ecosystems 31 requiring long-term monitoring.

Keywords: remote sensing, permafrost, larch forests, thermal anomaly, post-pyrogenic sites,
 seasonal thawed layer

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# 35 1. Introduction

36 Wildfire impact is the main factor, which affected strongly on the state of boreal ecosystems of 37 Siberia. The postfire changes in the vegetation cover in the larch forests of Central Siberia form 38 conditions for significant changes in thermal balance. These changes can affect the seasonal thawed 39 layer and its dynamics. Significant and long-term post-fire effects are well-documented in the 40 permafrost zone of Siberia [1-5]. A number of the problems associated with vegetation cover 41 disturbances have been discussed, such as degradation of seasonal thawed layer of permafrost as 42 well as variation of temperature and water regimes of soils [3,6–7]. Postfire changes in the thermal 43 balance can result in the disturbance of the "transitional layer", which protects the upper horizons of

44 permafrost [8]. Given the vast nature of the geographical area to be managed, satellite techniques

- 45 are the primary means for wildfire monitoring in most part of the boreal forest zone of Russia.
- 46 Evaluation of the effects of fires on Siberian ecosystems also requires the usage of remote data.

The main aims of study are (i) to perform a quantitative analysis of thermal anomalies in fire-damaged areas of the permafrost zone of Siberia and its dynamics according to the stages of post-fire succession, (ii) to obtain estimates of the depth of the seasonally thawed layer under

50 conditions of excessive heat flux on the surface based on numerical modeling technique.

# 51 2. Experiments

#### 52 2.1. Study area

The study area included the territory of Evenkia (central part of Siberia), covering the territory from 62° to 66° N and from 96° to 107° E (Fig. 1). This region belongs to the Central Siberian flat-mountainous taiga region of the boreal taiga zone. Siberian larch (*Larix sibirica, Larix gmelinii*) is the dominant species in the forest stands. The study area belongs to the continuous permafrost zone according to the Circum-Arctic permafrost and ground ice map by the National Snow and Ice Data Center [9]. Large-scale wildfires are typical for the territory, which are detected by satellite

59 monitoring only and are not served by fire protection system.



60	Figure 1. Study area in Central Siberia. Types of permafrost: 1 – continuous (90 – 100% of territory); 2
61	- discontinuous (50 - 90%); 3 - sporadic (10 - 50%); 4 - isolated patches (0 - 10%); 5 - remote sensing
62	monitoring data on burned area in 2016 – 2018.

#### 63 2.2. *Materials and method*

64 Retrospective multispectral materials of Terra and Aqua / MODIS for the period 2006–2018, as 65 well as information on wildfires in the format of geo-information polygonal layer [10] were used to 66 post-fire plots selecting. High resolution imagery (15-30 m) of Landsat/ETM/OLI (Enhanced 67 Thematic Mapper/ Operational Land Imager) and Sentilel-2 were used for wildfire geometry precise 68 estimating. Multispectral data from Terra/Aqua and the retrospective imagery allows evaluate 69 long-term changes using both the "vegetation" channels of the spectrum and the thermal range. The 70 characteristics of the postfire area were determined by analyzing of spectral features in the range of 71  $\lambda_1$  = 0.620–0.670 µm,  $\lambda_2$  = 0.841–0.876 µm (product MOD09GQ), and in thermal band of  $\lambda_3$  =

10.780–11.280 µm (product MOD11A1)/(L2G и L3 https://lpdaac.usgs.gov/dataset\_discovery/modis)
 [11,12].

Temperature, albedo and Normalized Difference Vegetation Index (NDVI) dynamics were investigated in total for 50 post-fire plots under different stages of succession (1, 2, 5 and 10 years after burning) in comparison with non-disturbed vegetation cover sites under the same natural conditions. Across the entire set of initial data, a 10-days averaging was performed taking into account the recovery succession stages. Dates of fires and post-fire stages were controlled using the

attribute information of the wildfire databank [10].

Variety of the thaw depth of the permafrost layer (Z) was estimated depending on the thermal
anomaly at the surface and vegetation recovery stage. Numerical modeling was based on the Stefan
solution of the heat conduction equation for the depth of the seasonal thawing layer [4,13,14]:

$$Z = \sqrt{\frac{2\lambda_f (T_s - T_f)\tau}{l \cdot u \cdot \rho}},$$
(1)

83 where  $\rho$  is the density (kg/m<sup>3</sup>), T is the temperature of the surface ( $T_s$ ) and the temperature of

84 permafrost layer ( $T_i$ ), x is the depth of the layer (m),  $\lambda$  is the thermal conductivity coefficient (W/(m

<sup>o</sup>C)) for the thawing layer ( $\lambda_1$ ) and permafrost layer ( $\lambda_2$ ),  $\tau$  is the duration of heating, *l* is specific heat

86 of fusion (J/kg), *u* is the volumetric water content of soil (%).

# 87 3. Results and discussion

88 Over the study area, fires were recorded on a total area of 12.743 MHa during 22 years of 89 available instrumental observations. It is more than  $\gamma > 12.0\%$  of forested area, relative burned area 90 (RBA) was  $0.51\pm0.18\%$  per year. Average RBA for Siberia is 1.19% [15]. For comparison, RBA is 91 0.56% for the forests of Western Canada [16]. Significant "accumulative" effect from fires provokes 92 large-scale anomalies of vegetation cover and thermal balance (Table 1).

**Table 1.** The averaged characteristics of post-pyrogenic sites in the mid-summer (maxima of thermal anomaly)

Time after burning, years	Anomalies of NDVI, %	Range of temperature maxima, °C	Anomalies of temperature, %
1	53.5±10.7	6.5-7.2	40–50
5	21.0±7.8	3.8–4.9	27–32
10	9.0±5.0	3.4-4.6	15-20

<sup>95</sup> 

96 On the postfire sites aged 1 year, the value of NDVI was typically of  $50 \pm 8\%$  comparing to the 97 non disturbed plots. The deviation of NDVI ( $21 \pm 7\%$ ) abnormality was 2 times lower of the control 98 values on the postfire sites aged 5 years. Fire sites with an age of 10 years do not differ significantly 99 from the control in terms of NDVI, which is caused by success restoration of on-ground vegetation 100 cover. The mean deviation from control values did not exceed 9% with a significant dispersion of  $\sigma =$ 101 5%. Anomalies in vegetation cover are lost during the next 5–7 years after fire. However, the process 102 of tree stands restoring is much longer up to 50 years [5,7,17].

For the same plots vegetation anomalies correlates with data on mid-summer thermal anomalies as well (Fig. 2, a). The maxima of temperature abnormality are staying significant more than 10–15 years after a wildfire (Table 1, Fig. 2). Absolute temperature maxima were fixed in range of  $7.0 \pm 1.5^{\circ}$ C during mid-summer for plots of burnt larch forest of Central Siberia. This is 20–40% higher than the temperature of undisturbed sites. The rate of loss of thermal anomalies was 2.5 times lower than the rate of vegetation restoration according to NDVI. The coefficients of exponential approximation are -0.08 and -0.2, respectively (Fig 2, b).



110 111 112

**Figure 2.** Experimental data on NDVI and temperature for post-fire plots relative to non-disturbed territories: (a) Thermal anomalies *vs* NDVI anomaly; (b) Long-term dynamics of NDVI index anomalies averaged for June – August (1) and averaged thermal anomalies (2).





117 Numeric simulation and field measurements [5,7,18] fix the average statistical norm of seasonal 118 thawing of active soil layer at 0.6-2.0 m under the conditions similar to the region of interest. As 119 shown by the results of numerical simulation, the excess heat flux on the surface causes an increase 120 in depth thawing (Z) on average by 10-20% (Fig. 3) relative to the average statistical norm. Thawing 121 of the soil profile is possible for an additional 0.5 m in depth under the conditions of stable 122 anticyclone that is usually observed over central part of Siberia during summer for up to 20-30 days 123 [19] and an excessive insolation of the surface. Despite the fact that the maximum of the thermal 124 anomaly is fixed in the middle of summer (Table 1), the maximum of thawing depth should be 125 expected in August. The delay in temperature dynamics with depth of soil profile should be taking 126 into account [20].

127 The results of satellite monitoring [4] and a number of ground-based experiments [7] allow us 128 to state the long-term (15–25 years) effects of thermal anomalies on the post-fire sites of the 129 permafrost zone. On post-fire plots the NDVI values are restored 3 to 5 years after the fire. But 130 surface temperature anomalies still can be considered as statistically significant during 10–15 years

- 131 after burning. The low rate of vanishing of the thermal anomaly makes it possible to consider this
- 132 factor of long-term influence on the state of the soil and the permafrost as one of the most significant.
- 133 A similar effect was observed, in particular, in Alaska [6]. The obtained results are also consistent
- 134 with the data for Russian forests [3,21]. This should be taken into account when assessing the state of
- 135 the forests, tree mortality or monitoring of regeneration processes in the forests of the permafrost
- 136 zone using satellite methods [22]. Temperature fields restored using the satellite data can be
- 137 considered as the basis for monitoring the condition of forests of the permafrost zone of Siberia.
- 138 Considering the predicted climate changes, toughening of fire regimes and increased fire
- activity in the forests of the permafrost zone [15,19,23,24], the integral effect of post-fire thermal
- 140 anomalies will strengthen and is likely to grow increasingly.

# 141 4. Conclusions

142 Changes in thermal regime of postfire areas in Central Siberian are accompanied by an 143 abnormal increase in average temperatures of the soil surface by  $\Delta T = 7.2 \pm 1.3$ °C. This is 20–40% 144 higher than the temperature of undisturbed sites. The NDVI values are restored 3 to 5 years after the 145 fire. The rate is depending on the regeneration of the vegetation cover. The partial regeneration of 146 vegetation covers does not compensate the changes, which lead to long-term temperature 147 abnormalities. The thermal balance of postfire sites with disturbed vegetation cover remain affected 148 for more than 15 years.

- 149 It was found that the abnormal temperatures on a significant area of the permafrost zone can 150 result in a seasonal increase in the thawing depth of the soil by 10–20% when compared with the 151 mean normal value.
- Vast postfire disturbances (currently up to 25% of total forested area per the last two decades) of the vegetation cover in the northern regions of Siberia have a significant effect on the temperature regime of the "ground cover"–"soil"–"permafrost layer" boundary layer. A more detailed study of post-fire effects is important for predicting the response of boreal ecosystems to the fire impact, which currently tends to increase.

157 The low rate of the thermal anomaly lost, at least in the first 10 years after the fire, makes it 158 possible to consider this factor of long-term influence on the state of the seasonal thawed soil layer as 159 one of the most significant. This technique and remote sensing data could be used for determining 160 the stability of permafrost ecosystems.

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  165 tools; E.P. wrote the paper.
- 166 **Conflicts of Interest:** The authors declare no conflict of interest.

# 167 Abbreviations

- 168 The following abbreviations are used in this manuscript:
- 169 MODIS: Moderate Resolution Imaging Spectroradiometer
- 170 NDVI: Normalized Difference Vegetation Index
- 171 RBA: Relative Burned Area
- 172 ETM: Enhanced Thematic Mapper
- 173 OLI: Operational Land Imager

# 174 References

175 1. Ivanova, G.A. Extreme fire-endangered seasons in forests of Evenkia. *Siber. Ecolog. J.* **1996**, *3*, *1*, 29–34. (in Russian)

- 177 2. Sofronov, M.A.; Volokitina, A.V.; Kajimoto, T. Ecology of wildland fires and permafrost: their
  178 interdependence in the northern part of Siberia. Proc. of 8-th Symp. on the Joint Sib. Permafr. Stud. 1999,
  179 211–218.
- 180 3. Anisimov, O.A.; Sherstiukov, A.B. Evaluating the effect of environmental factors on permafrost in Russia.
   181 *Kriosfera Zemli (Earth's Cryosph).* 2016, *XX*, *2*, 90–99. (in Russian)
- Ponomarev, E.I.; Ponomareva, T.V. The Effect of Postfire Temperature Anomalies on Seasonal Soil Thawing in the Permafrost Zone of Central Siberia Evaluated Using Remote Data. *Contemp. Probl. of Ecol.* **2018**, *11*, *4*, 420–427. DOI:10.1134/S1995425518040066.
- 185 5. Knorre, A.A.; Kirdyanov, A.V.; Prokushkin, A.S.; Krusic, P.J.; Buntgen, U. Tree ring-based reconstruction
  186 of the long-terminfluence of wildfires on permafrost active layer dynamics in Central Siberia. *Science of the*187 *Total Env.* 2019, 652, 314–319. doi:https://doi.org/10.1016/j.scitotenv.2018.10.124.
- Brown, D.R.N.; Jorgenson, M.T.; Kielland, K.; Verbyla, D.L.; Prakash, A.; Koch, J.C. Landscape Effects of
  Wildfire on Permafrost Distribution in Interior Alaska Derived from Remote Sensing. *Remote Sens.* 2016, *8*, *8* (654), 22. DOI:10.3390/rs8080654.
- 191 7. Bezkorovaynaya, I.N.; Borisova, I.V.; Klimchenko, A.V.; Shabalina, O.M.; Zakharchenko, L.P.; Il'in, A.A.;
  192 Beskrovny, A.K. Influence of the pyrogenic factor on the biological activity of the soil under permafrost
  193 conditions (Central Evenkia). *Vestnik KrasGAU*. 2017, *9*, 181–189. (in Russian)
- 194 8. Desyatkin, R.V.; Desyatkin, A.R. Temperature Regime of Solonetzic Meadow-Chernozemic
  195 Permafrost-Affected Soil in a Long-Term Cycle. *Eurasian Soil Sc.* 2017, 50, 1301–1310. DOI:
  196 10.1134/S1064229317090022.
- Brown, J.; Ferrians, O.J.; Heginbottom, J.A.; Melnikov, E.S. Circum-arctic map of permafrost and ground ice conditions, Ver. 2. 2002. Boulder, Colorado USA: National Snow and Ice Data Center. Digital media.
   [https://nsidc.org/data/ggd318].
- Ponomarev, E.I.; Shvetsov, E.G. Satellite survey of forest fires and GIS methods for data alignment. *Issled. Zemli Kosmosa* (Rem. Sens.) 2015, *1*, 84–91. DOI: 10.7868/S0205961415010054. (in Russian)
- 202 11. Wan, Z.; Hook, S.; Hulley, G. MOD11A1 MODIS/Terra Land Surface Temperature/Emissivity Daily L3 203 V006 [Data EOSDIS LP DAAC. 2015. Global 1km SIN Grid set]. NASA 204 DOI:10.5067/MODIS/MOD11A1.006.
- 205 12. Vermote, E.; Wolfe, R. MOD09GQ MODIS/Terra Surface Reflectance Daily L2G Global 250m SIN Grid
   206 V006 [Data set]. NASA EOSDIS LP DAAC. 2015. DOI:10.5067/MODIS/MOD09GQ.006.
- Arzhanov, M.M.; Eliseev, A.V.; Demchenko, P.F.; Mokhov, I.I. Simulation of dynamics of temperature and hydrological regime of near-surface permafrost using climatic data (reanalysis). *Kriosfera Zemli (Earth Cryosphere)*. 2007, 11, 4, 65–69. (in Russian)
- 210 14. Vinogradov, Yu.B.; Semenova, O.M.; Vinogradova, T.A. Hydrological modeling: calculation of the
  211 dynamics of thermal energy in the soil profile. *Kriosfera Zemli (Earth Cryosphere)*. 2015, 19, 1, 11–21. (in
  212 Russian)
- 213 15. Ponomarev, E.I.; Kharuk, V.I.; Ranson, J.K. Wildfires Dynamics in Siberian Larch Forests. *Forests* 2016, 7, 125, doi:10.3390/f7060125.
- 215 16. de Groot, W.J.; Cantin, A.S.; Flannigan, M. D.; Soja, A. J.; Gowman, L. M.; Newbery, A. A comparison of
  216 Canadian and Russian boreal forest fire regimes. *For. Ecol. and Manage.* 2013, 294, 23–34.
  217 DOI:10.1016/j.foreco.2012.07.033.
- Kharuk, V.I.; Dvinskaya, M.L.; Ranson, K.J. Spatio-temporal dynamics of fires in the larch forests of the northern taiga of Central Siberia. *Rus. J. of Ecol.* 2005, *5*, 1–10. (in Russian)
- 18. Abaimov, A.P.; Sofronov M.A. The main trends of post-fire succession in near-tundra forests of central
  Siberia. In *Fire in ecosystems of boreal Eurasia*, Goldammer, J.G., Furyaev V.V., Eds.; Dordrecht: Kluwer
  Academic Publishers, 1996; pp. 372–386.
- Valendik, E.N.; Kisilyakhov, E.K.; Ryzhkova, V.A.; Ponomarev, E.I.; Danilova, I.V. Conflagration fires in taiga landscapes of Central Siberia. *Geogr. and Natur. Resour.* 2014, 35, 1, 41–47. DOI 10.1134/S1875372814010065.
- 20. Desyatkin, R.V.; Desyatkin, A.R.; Fedorov, P.P. Temperature regime of taiga cryosoils of Central Yakutia.
   227 *Kriosfera Zemli (Earth Cryosphere).* 2012, *16*, *2*, 70–78. (in Russian)
- 228 21. Lebedeva, L.S.; Semenova, O.M.; Vinogradova, T.A. Calculation of the depth of seasonally melted layer in
   various landscapes of Kolyma water-balance station based on Gidrograf hydrological model. *Kriosfera*

- 230 Zemli (Earth Cryosphere). 2015, 19, 2, 35–44. (in Russian)
- 231 22. Bartalev, S.A.; Stytsenko, F.V.; Egorov, V.A.; Loupian, E.A. Satellite-Based Assessment of Russian Forest
   232 Fire Mortality. *Lesovedenie (Forestry)*. 2015, *2*, 83–94. (in Russian)
- 23. Vaganov, E.A.; Arbatskaya, M.K. The history of climate and fire frequency in the central part of
  234 Krasnoyarsk krai. 1. Climate conditions of a season and distribution of fires during a season. *Siber. Ecolog.*235 J. 1996, *3*, *1*, 9–18. (in Russian)
- 236 24. Flannigan, M.; Stocks, B.; Turetsky, M.; Wotton, M. Impacts of climate change on fire activity and fire 237 circumboreal Biol. management in the forest. Glob. Change 2009, 15, 49-560. 238 doi:10.1111/j.1365-2486.2008.01660.x.
- 239 25. Kharuk, V.I.; Dvinskaya, M.L.; Petrov, I.A.; Im, S.T.; Ranson, K.J. Larch Forests of Middle Siberia:
  240 Long-Term Trends in Fire Return Intervals. *Reg. Environ. Chang.* 2016, *8*, 2389–2397.



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